## Long-period ground motion simulation of great Nankai Trough, Japan, earthquakes

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### The Nankai trough earthquakes



Central Disaster Prevention Council (2011)

## Long period ground motion

- The long-period ground motions cause damage to highrise and large-scale structures.
  - 1985 Michoacan, Mexico earthquake M8.0
  - 2003 Tokachi-oki, Japan earthquake (Mw8.3)
  - 2011 Tohoku, Japan earthquake (Mw9.0)



(USGS)



(Sapporo Fire Bureau)



(youtube)

## Long-Period Ground Motion Hazard Maps

「長周期地震動予測地図」

- It is difficult to determine the source model of the anticipated earthquake.
- For improving seismic hazard assessment to prepare for the anticipated Nankai Trough earthquake, it is important to understand a possible range of ground motion caused by uncertainty of source model.

平成 21 年(2009 年) 地震調査研究推進本部 地震調査委員会 地震調査研究推進全部地震調査委員会

Earthquake Research Committee (2009)

Earthquake Research Committee (2012)

## Objective

- Simulate long-period ground motions for anticipated Nankai Trough earthquake (M8-9) using various source models and analyze a variability of predicted long-period ground motions.
- For selecting source models,
  - Assuming various source parameters: rupture area, asperity (area of large slip) location, and hypocenter.
  - Assuming the large slip near the trough following the lesson from the 2011 Tohoku earthquake.
- The long-period ground motions are simulated by the finite difference method using characterized source model and recently developed three-dimensional velocity structure model.

## Characterized source model

- The characterized source model is used in the long-period ground motion simulation.
- Three kinds of parameters
  - Outer source parameters
    - Shape and area of the whole fault
  - Inner source parameters
    - Heterogeneity of the source fault
  - Other source parameters
    - Rupture process





# Flowchart of determination of source parameters for subduction-zone interplate earthquakes



Earthquake Research Committee (2009); Fujiwara et al. (2009)

# Characterized source model for the Nankai trough earthquakes





- Source areas (14 cases)
  - Single-segment earthquake
    - Nankai (ANNKI: Mw 8.5), Tonankai (ATNKI: Mw 8.2), Tokai (ATOKI: Mw 8.0), Hyuga-nada (AHGND: Mw 8.3), and along the trough (ATRGH: Mw 8.1)
    - Along the trough area has 3 cases (a, b, c).
  - Multi-segment earthquake (simultaneous rupture of multi segment)
  - 2-segment rupture
    - Nankai + Tonankai (ANNI1: Mw 8.7)
    - Tonankai + Tokai (ANNI2: Mw 8.4)
  - 3-segment rupture
    - Nankai + Tonankai + Tokai (ANNI3: Mw 8.8)
  - 4-segment rupture
    - 3-segment + along the trough (ANNI4: Mw 8.9)
    - 3-segment + Hyuga-nada (ANNI5: Mw 8.9)
  - Seismic moment are calculated using scaling model

# Characterized source model for the Nankai trough earthquakes



- Hypocenter (3 cases)
  - West: boundary of Hyuda-nada and Nankai
  - Center: boundary of Nankai and Tonankai
  - East: boundary of Tonankai and Tokai
- Asperity (2 cases)
  - Deeper and shallower cases
  - Total area of asperities is 20 % of source fault area (Murotani et al., 2008)
  - The number of asperity is one for along the trough, and three (area ratio is 2:1:1) for Nankai, Tonankai, Tokai and Hyuga-nada.



# Characterized source model for the Nankai trough earthquakes



- Source time function
  - Kostrov-like slip velocity time function : Nakamura

& Miyatake (2000)

- Rupture velocity
  - 2.7 km/s





## Velocity structure model

- Vp(km/s) Vs(km/s)  $\rho$  (g/cm<sup>3</sup>) Qp Qs Layer 70 0.35 1.80 119 1 1.7 2 1.8 0.5 1.95 170 100 3 2.0 2.00 204 120 0.6 4 2.1 2.05 238 140 0.7 5 2.2 0.8 2.07 272 160 180 6 2.3 0.9 2.10 306 7 2.4 1.0 2.15 340 200 8 2.7 1.3 2.20 442 260 300 9 3.0 2.25 510 1.5 10 3.2 2.30 578 340 1.7 11 3.5 2.0 2.35 680 400 400 12 4.2 2.4 2.45 680 13 5.0 2.9 400 2.60 680 400 14 5.5 3.2 680 2.65 15 400 680 5.8 3.4 2.70 400 16 6.4 3.8 2.80 680 500 17 7.5 4.5 3.20 850 18 5.0 2.9 2.40 340 200 6.8 19 4.0 2.90 510 300 500 20 8.0 4.7 3.20 850 21 5.4 2.8 2.60 340 200 22 6.5 3.5 2.80 510 300 23 8.1 4.6 3.40 850 500
- Japan Integrated Velocity Structure Model

7<sup>th</sup> layer lower surface : sedimentary wedge



17<sup>th</sup> layer upper surface : Moho discontinuity



18<sup>th</sup> layer upper surface : Philippine Sea plate





Earthquake Research Committee (2012)

## Finite difference method

- 3-D finite difference (FD) simulation with discontinuous grids (Aoi and Fujiwara, 1999)
- Grid spacing
  - 200m in horizontal, 100m in vertical (shallower than 8km)
  - 600m in horizontal, 300m in vertical (8km 70km)
- Area of the 3-D simulation
  - 1100 km x 700 km x 70 km (The total number of grid points = 2 billion)
- Inelastic attenuation (Graves, 1996) : reference period of 5 sec
- Valid period of the FD simulation : more than about 5 second
  - Since the characterized source model does not include short-wavelength heterogeneity.
- Computation time
  - 36 hours/ 1 scenario
    - 60000 steps (120 Hz)
    - Itanium 1.66GHz × 256Core
    - Memory 130 GB

## Velocity waveforms @ OIT

## Simultaneous rupture of 3 and 4 segments (30 scenarios)





## Velocity waveforms @ OSK

## Simultaneous rupture of 3 and 4 segments (30 scenarios)





### Velocity waveforms @ AIC Simultaneous rupture of 3 and 4 segments (30 scenarios)





## Velocity waveforms @ TKY

## Simultaneous rupture of 3 and 4 segments (30 scenarios)

TKY Velocity (cm/s)





Long duration for west-hypocenter scenarios

## Character of the waveforms



ΟΙΤ	<ul> <li>Smaller amplitude for <u>west-hypocenter</u> scenarios</li> <li>S waves from deeper asperities of the Nankai area</li> <li>Later phases from the along trough 'a' area</li> </ul>
OSK	<ul> <li>Larger amplitude for <u>west-hypocenter</u> scenarios</li> <li>Later phases from the along trough 'a' area</li> </ul>
AIC	<ul> <li>Larger amplitude for <u>west-hypocenter</u> scenarios</li> <li>Later phases from the along trough 'c' area</li> </ul>
ТКҮ	<ul> <li>Long duration for west-hypocenter scenarios</li> </ul>

The hypocenter location tends to have a greater impact on long-period ground motions than the asperities location and the rupture areas.

## Velocity response (damping factor of 5%)



# Frequency distribution of Peak ground velocity (PGV) and Velocity response (Sv)



## Maps of Peak ground velocity (PGV) and Velocity response (Sv)

Maximum



Median the generation of long-period ground motions



20

50

50 100 200 300

100

200

PGV

Interquartile range (IQR)







Long-period ground motions in the Kanto basin are relatively large for most cases.





Long-period ground motions in the Osaka and Nobi basins are greatly influenced by a few specific scenarios.

### Maps of PGV and Sv (without 4-segment scenarios)

Maximum









#### Median



### Interquartile range (IQR)











## Maps of PGV and Sv (before and after applying new findings)

Maximum

Maximum



#### Median



#### Interquartile range





- Hyuga-nada
- Along the trough



Median



#### Interquartile range



## Effect of slip velocity function

- Change the slip velocity function for along the trough area
  - Kostrov-like slip velocity time function (Nakamura and Miyatake, 2000)
  - Boxcar-like slip velocity time function



PGV ratio (Boxcar / Kostrov)

PGV values decrease by a factor of two locally.
 Selection of appropriate source time function is important.

## Summaries

- We simulated long-period ground motion of the Nankai trough earthquakes using 55 scenarios with various possible source parameters.
- The hypocenter location (rupture start point) tends to have a greater impact on long-period ground motions than the asperities and the rupture areas.
- Long-period ground motions in the Osaka and Nobi basins are greatly influenced by a few specific scenarios including the along trough area.
- Long-period ground motions in the Kanto basin are relatively large for most cases.
- New findings will change the possible range of longperiod ground motion.

### Future tasks

- Selection of appropriate source time function for the along trough area.
- Improvement of source model, velocity structure model, and simulation method.
  - Introducing some kind of heterogeneity into the characterized source model.
  - Modeling shallow lower velocity layers.

– etc...