



# Ground-Motion Prediction Equation ( GMPE) for Taiwan

Po-Shen Lin



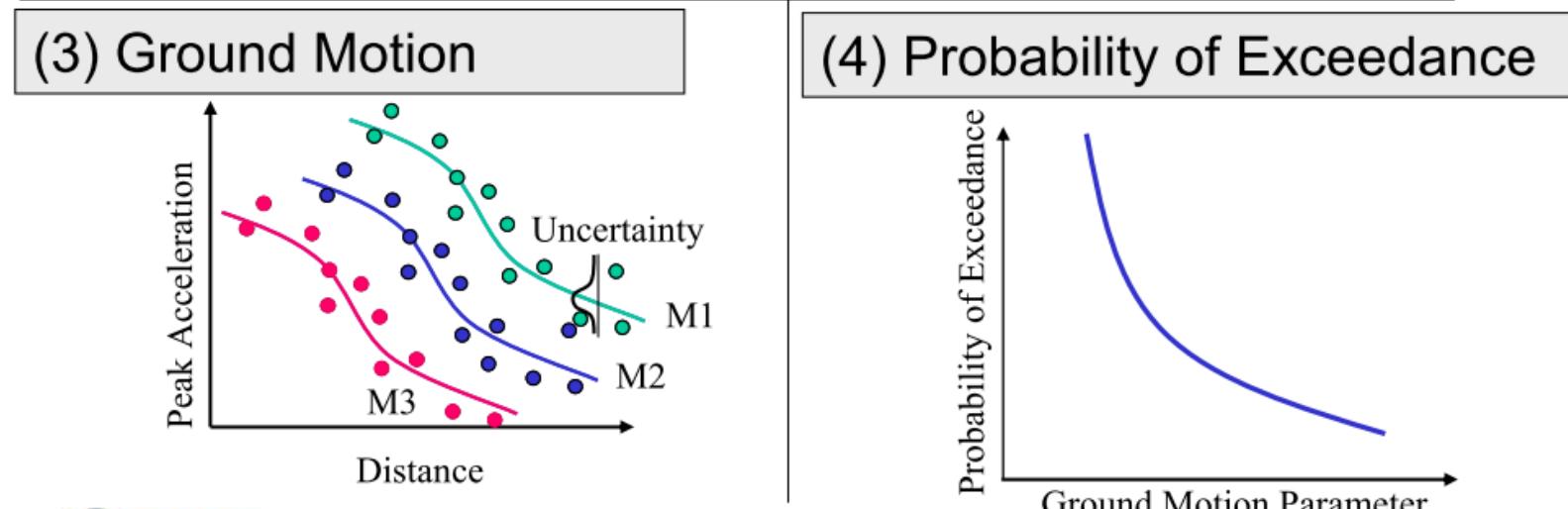
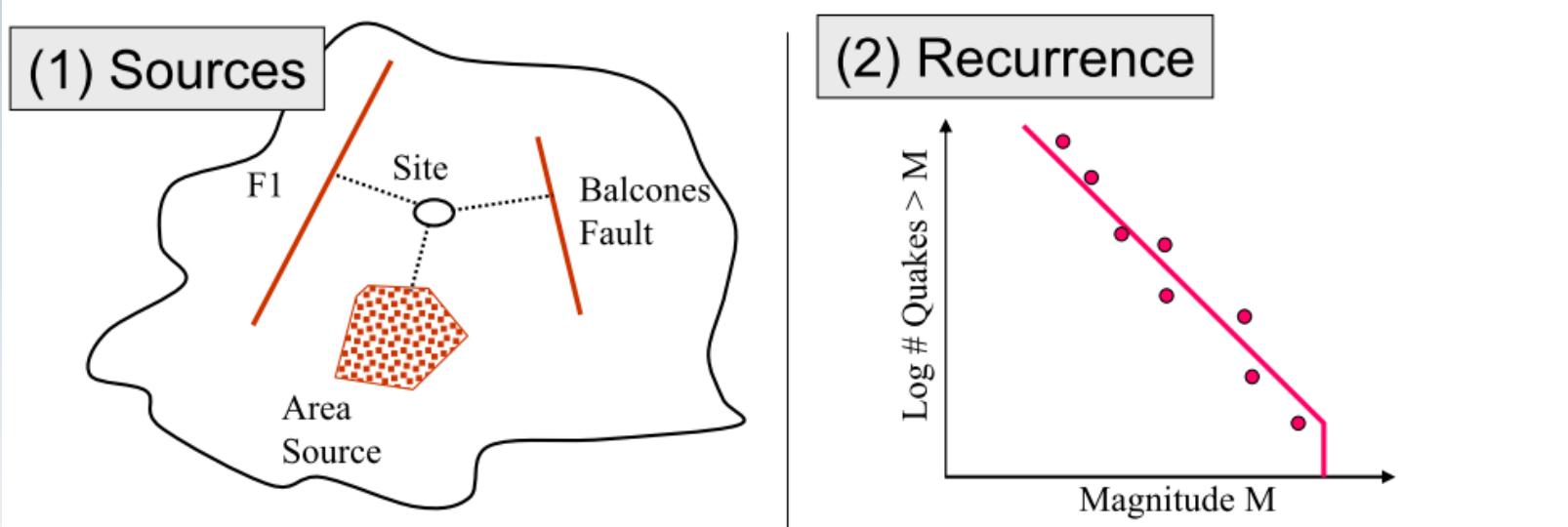
Disaster Prevention Technology Research Center,  
Sinotech Engineering Consultants, INC.



## Outline

- Introduction of GMPE
  - Important component of PSHA
  - From attenuation equation to GMPE
  - Steps for building GMPE
- Taiwan's GMPE – Past 、 Present 、 Future
- Conclusion
- Future work

## Steps in Probabilistic Seismic Hazard Analysis



FEMA

Instructional Material Complementing FEMA 451, Design Examples

Seismic Hazard Analysis 5a - 24



## What is GMPE

- “Was”
  - Attenuation relations
  - Attenuation relationships
  - Attenuation equations
- It is an equation that can be used to predict the possible ground-motion value during a future earthquake.
- Most of them are “empirical”, and was developed from a set of ground-motion data with proper physical meaning



## Why Call them “Ground-Motion Prediction Equations”

- “**Attenuation Equations**” is a poor term
  - They describe the **INCREASE** of amplitude with magnitude at a given distance
  - They describe the **CHANGE** of amplitude with distance for a given magnitude (usually, but not necessarily, a **DECREASE** of amplitude with increasing distance).

Art McGarr, 2006



## Ground Motion Prediction Equations

- Empirical regressions of recorded data
- Estimate ground shaking parameter (peak ground acceleration, peak velocity, spectral acceleration or velocity response) as a function of
  - (1) magnitude
  - (2) distance
  - (3) site
- May consider fault type (strike-slip, normal, reverse)

Art McGarr, 2006



## Steps for building GMPE

- Establish database
- Select form of predictive equation
- Perform regression analyses
- Evaluate uncertainty



## Combination of horizontal measurements

- Arithmetic mean (算數平均數)
- Both
- Geometric mean (幾何平均數)
- Largest component
- Random
- Resultant ( $a=a_1*\cos\theta+a_2*\sin\theta$ )
- Vectorial addition (  $a_v = \sqrt{a_1(t)^2 + a_2(t)^2}$  )
- GMRotI50



## Characterisation of source

- Earthquake magnitude,  $M$ 
  - $M_L$
  - $M_S$
  - $m_b$
  - $M_W$
- Source mechanism
  - Strike slip 、 normal faulting 、 reverse
- Tectonic setting
  - Crustal 、 subduction



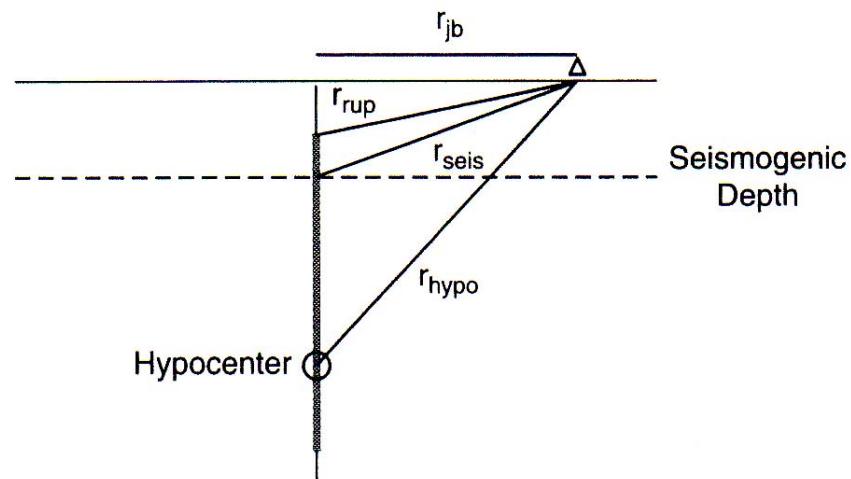
## Characterisation of path

- Definitions of source-to-site distance
  - Epicentral distance 、 hypocentral distance 、 rupture centroid distance 、 centre-of-energy-release distance 、 surface projection distance ( $R_{jb}$ ) 、 surface projection distance with focal depth 、 rupture distance ( $R_{rup}$ ) 、 seismogenic distance 、 average site to rupture end distance 、 equivalent hypocentral distance (EHD)

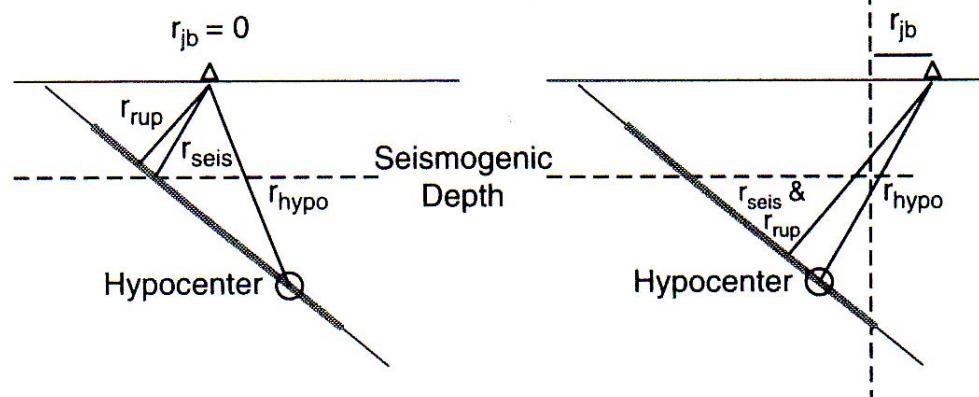


## source-to-site distance

Vertical Faults



Dipping Faults



(Abrahamson and Shedlock, 1997)



- From data selection
- Multiplicative factor
- Shear-wave velocity (Vs30)



## Form of the predictive equation

$$Y=f(M, R, P_i)$$

$Y$ : Ground motion parameter of interest

- $M$ : The magnitude of the earthquake
- $R$ : Distance from the source to the particular site
- $P_i$ : Other parameters (earthquake source, local site conditions, wave propagation path...)



## Analysis techniques

- The majority of ground motion estimation studies use the *ordinary least squares method*
- Two-stage method
- Non-linear regression
- Maximum-likelihood method (mixed effects method)



Author	Equation	Parameter	Region	Data	Events	Records
茅聲燾 (1978)	$Y = 0.3725e^{0.876M} D^{1.836}, D=(R^2+400)^{1/2}$ (g)	$R$ 為震源距離	台灣地區	CWB震度	5	
Tsai & Bolt (1983)	$PGA = 17.5e^{0.869M}(R+0.0606e^{0.700M})^{-1.09}$ (gal)	$R$ 為震源距離， $M$ 為芮氏規模( $M_L$ )	台灣東北部	SMART-1	6	
倪與邱 (1991)	$\ln(PGA) = 6.09 + 0.26M - 0.87\ln(R+6.9)$ (gal)	$R$ 為震央距離， $M$ 為芮氏規模( $M_L$ )	台灣地區 (岩盤)	SMA-1	49	
Chiu and Ni (1993)	$\ln(PGA) = 4.15 + 1.41M_L - 2.37\ln(R+13.7)$ (gal)	$R$ 為震央距離， $M$ 為芮氏規模( $M_L$ )	花蓮地區	SMART-2		
黃正耀 (1995)	$Y = 0.0253e^{1.5873M}(R+0.3155e^{0.6165M})^{-2.3027}$ (g)	$R$ 為測站至斷層破裂面之最短距離， $M$ 為芮氏規模( $M_L$ )	台灣地區	TSMIP (7個地震)	7	526
劉坤松 (1996)	$Y = 0.0308e^{1.20M}(R+0.1413e^{0.6892M})^{-1.741}$ (g)	$R$ 為震源距離， $M$ 為芮氏規模( $M_L$ )	台灣東北部地區	TSMIP		
羅俊雄 (1996)	$Y = 0.0267e^{1.354M}(R+0.2138e^{0.7499M})^{-1.0329}$ (g)	$R$ 為震源距離， $M$ 為芮氏規模( $M_L$ )	台灣地區 (岩盤)	SMA-1		
羅俊雄 (1996)	$Y = 0.0273e^{0.1058M}(R+0.141e^{0.656M})^{-1.6472}$ (g)	$R$ 為測站至斷層破裂面之最短距離， $M$ 為芮氏規模( $M_L$ )	台灣地區	TSMIP		
辛在勤 (1998)	$PGA = 12.44e^{1.31M_L} r^{1.487}$ (gal)	$r$ 為震源距離， $M_L$ 為芮氏規模( $M_L$ )	台灣地區	TSMIP	22	
劉坤松 (1999)	$\ln(PHA) = -1.339\ln(r+2.12) - 0.0071r + 1.167M + 2.192 + G(1.475\ln(r) - 6.792)$ $\ln(PVA) = -1.681\ln(r+2.45) - 0.0036r + 1.250M + 2.210 + G(1.009\ln(r) - 4.464)$ $\ln(PHV) = -1.377\ln(r+1.67) - 0.0023r + 1.581M - 3.070 + G(1.066\ln(r) - 4.909)$ $\ln(PVV) = -1.415\ln(r+1.58) - 0.0012r + 1.666M -$	$r$ 為震源距， $M$ 為地震矩規模	台灣地區	TSMIP	35	2187



## Most used GMPE for the Engineer

Campbell's form

$$a = b_1 e^{b_2 M} (R + b_4 e^{b_5 M})^{-b_3}$$

Joyner and Boore's form

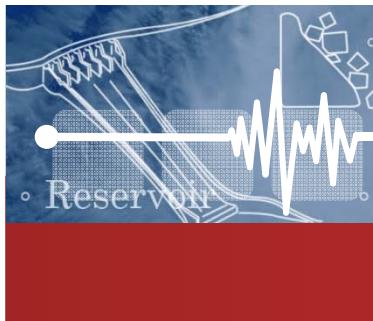
$$\log_{10}(a) = b_1 + b_2 M + b_3 \log_{10} \left[ (R^2 + b_5^2)^{b_4} \right]$$

Kanai's form

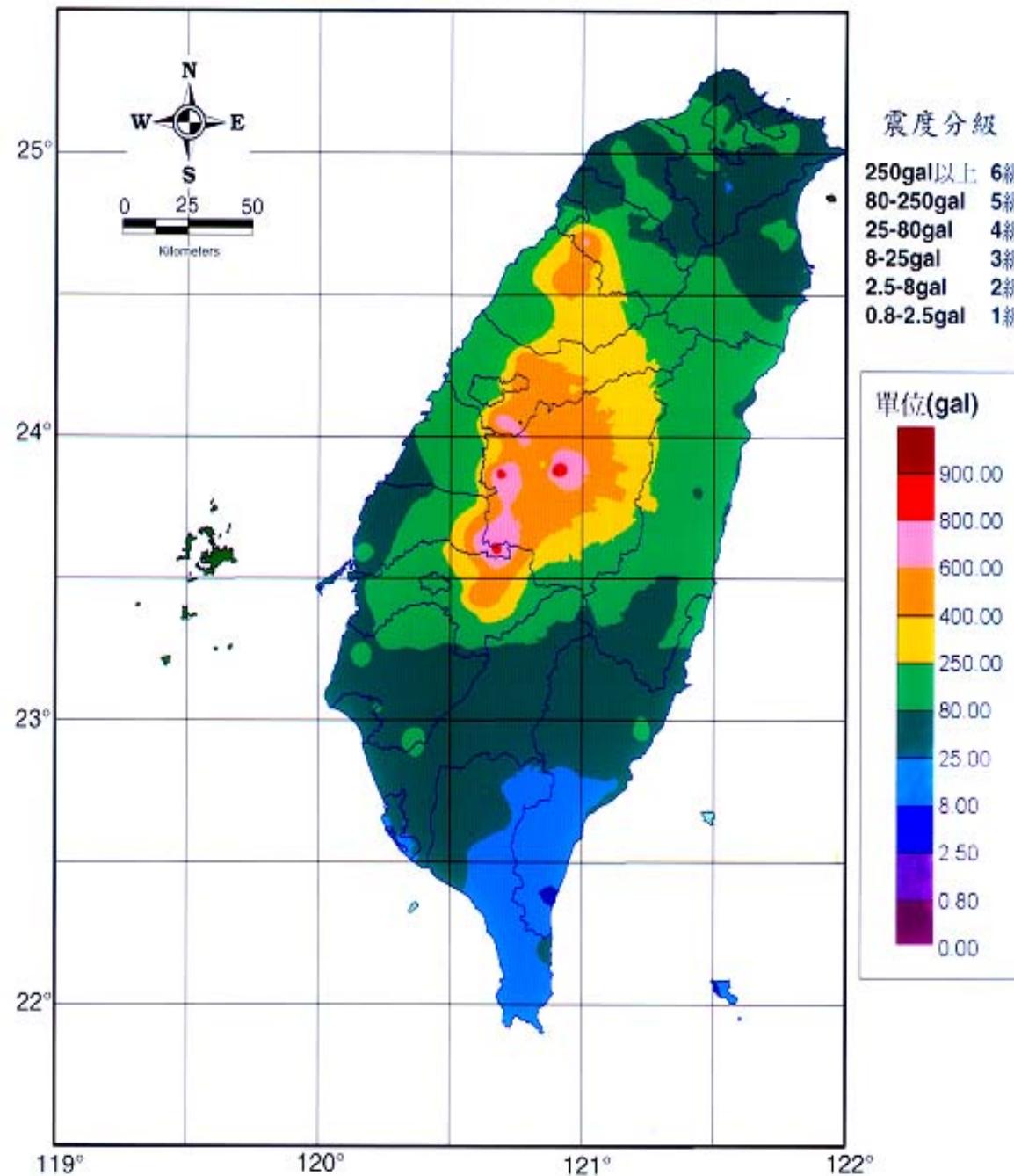
$$a = b_1 e^{b_2 M} (R + b_4)^{-b_3}$$

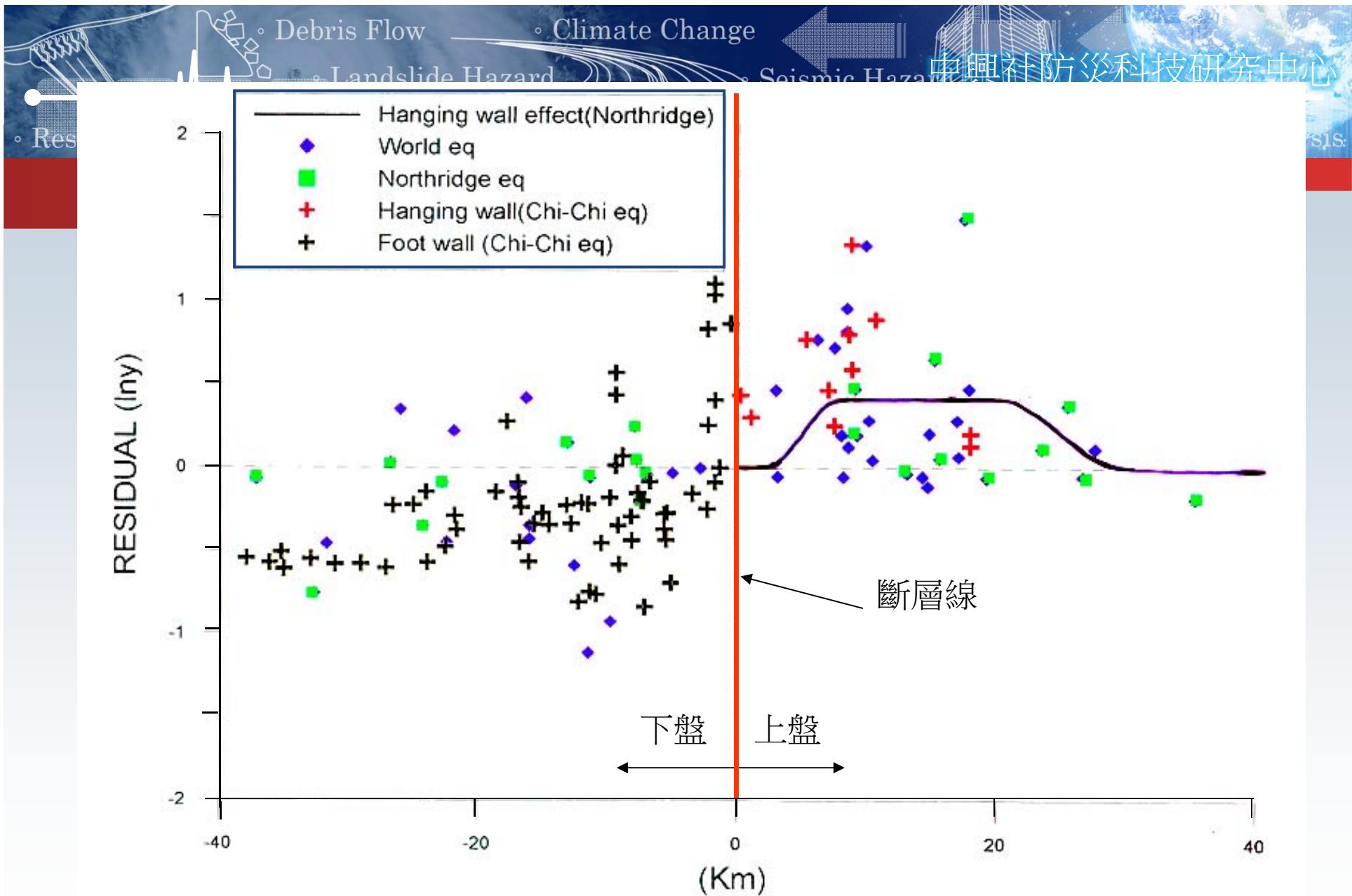
Japan Rock Site's form

$$\log_{10}(0.981a) = \left( \frac{R + b_4}{b_5} \right) \cdot (-b_1 + b_2 M - b_3 M^2)$$



## 921 集集地震台灣地區震度圖

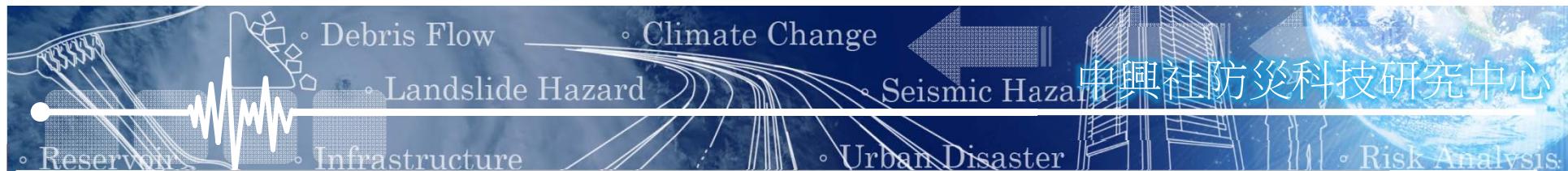




集集大地震與 Northridge 地震及世界其他大地震比較上盤效應



李錫堤等 (1999)			台灣地區 美濃水庫	TSMIP		
Loh et al. (2000)	$Y = 0.02968e^{1.2M} (R + 0.1464e^{0.698M})^{-1.7348}$ (g)	$R$ 為測站至斷層破裂面之最短距離， $M$ 為( $M_W$ )？	台灣地區	TSMIP		
羅俊雄與 溫國樑， 2000	Spectral Acceleration at 80 periods (0.029~10 sec)		台灣地區	TSMIP		
Chang et al. (2001)	$\ln A = 2.8096 + 0.8993M - 0.4381 \ln D_p -$ $(1.0954 - 0.0079 D_p) \ln D_e$ $\ln A = 4.7151 + 0.8468M - 0.1745 \ln D_p - 1.2972$ $\ln D_h$	$D_e$ 、 $D_h$ 、 $D_p$ 分別為震 央距、震源距與震 源深度(km)	台灣地區	TSMIP	45 19	
趙曉玲 (2001)			台灣地區	TSMIP	9	3104
Wu et al. (2001)	$\log_{10}(\text{PGA}) = 0.00215 + 0.581M - \log_{10}(r_{rup} + 0.008$ $71 \times 100.5M) - 0.00414r_{rup}$ $\log_{10}(\text{PGV}) = -2.49 + 0.810M - \log_{10}(r_{rup} + 0.00871 \times$ $100.5M) - 0.00268r_{rup}$	$M$ 為地震矩規模 $M_W$ ， $r_{rup}$ 為距斷層破 裂面距離(km)	台灣地區	TSMIP	60	1941



Jean(2001)	$Y = 0.00369e^{1.75377M} (R + 0.12220e^{0.78315M})^{-2.05644}$	震央距以及距斷層面最短距離	台灣地區	TSMIP			
Liu and Tsai (2005)			台灣地區	TSMIP	51	7900	
Jean et al. (2006)	$PGA = 0.0028e^{1.7331M} (R + 0.0999e^{0.7719M})^{-2.0639}$ $Sa03 = 0.0079e^{1.7253M} (R + 0.1199e^{0.7850M})^{-2.0489}$ $Sa10 = 0.0027e^{1.7731M} (R + 0.1154e^{0.7714M})^{-2.0419}$	堅硬地盤, $M_w$ , 震央距以及距斷層面最短距離	台灣地區	TSMIP			
吳芳儒 (2008)	$\ln Y_{ij} = C_1 + C_2 M_i + C_3 (R_{ij} + C_5 10^{C_4 M_i}) + C_4 \ln(R_{ij} + C_5 10^{C_4 M_i})$ $\ln Y_{ij} = C_1 + C_2 M_i + C_3 (R_{ij} + C_5 10^{C_4 M_i}) + C_4 \ln(R_{ij} + C_5 10^{C_4 M_i}) + C_5 \ln(V_{s30})$ $\ln Y_{ij} = C_1 + C_2 M_i + C_3 (R_{ij} + C_5 10^{C_4 M_i}) + C_4 \ln(R_{ij} + C_5 10^{C_4 M_i}) + C_5 \ln(A_j)$	地震矩規模 ( $M_w$ ) 震央距	台灣地區 (2006屏東地震)	TSMIP	3 (186)	921 (5852)	
Lin and Lee (2008)	Subduction zone earthquake	地震矩規模 $M_w$ 距斷層面最短距離 ( $R_{rup}$ )	台灣地區 (隱沒帶)	TSMIP	54	4823	
Lin et al.(2011)	Crustal earthquake (hanging-wall, footwall)	地震矩規模 $M_w$ 距斷層面最短距離 ( $R_{rup}$ )	台灣地區 (地殼地震)	TSMIP	52	5268	
Lee et al. (2012)	Arias Intensity	地震矩規模 $M_w$ 距斷層面最短距離 ( $R_{rup}$ )	台灣地區	TSMIP	62	6570	

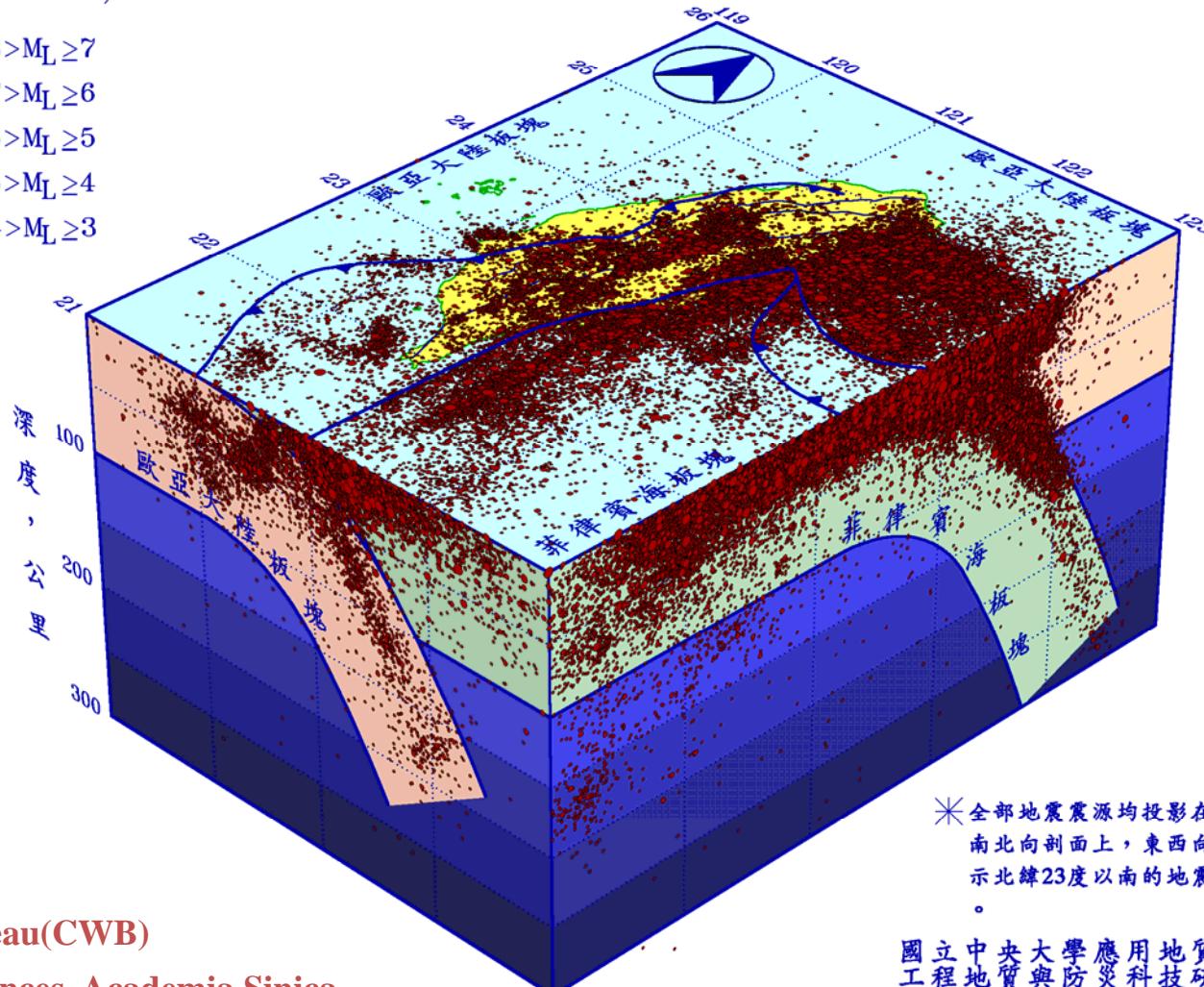


## Seismotectonic framework of Taiwan

### 臺灣的地震與地體構造

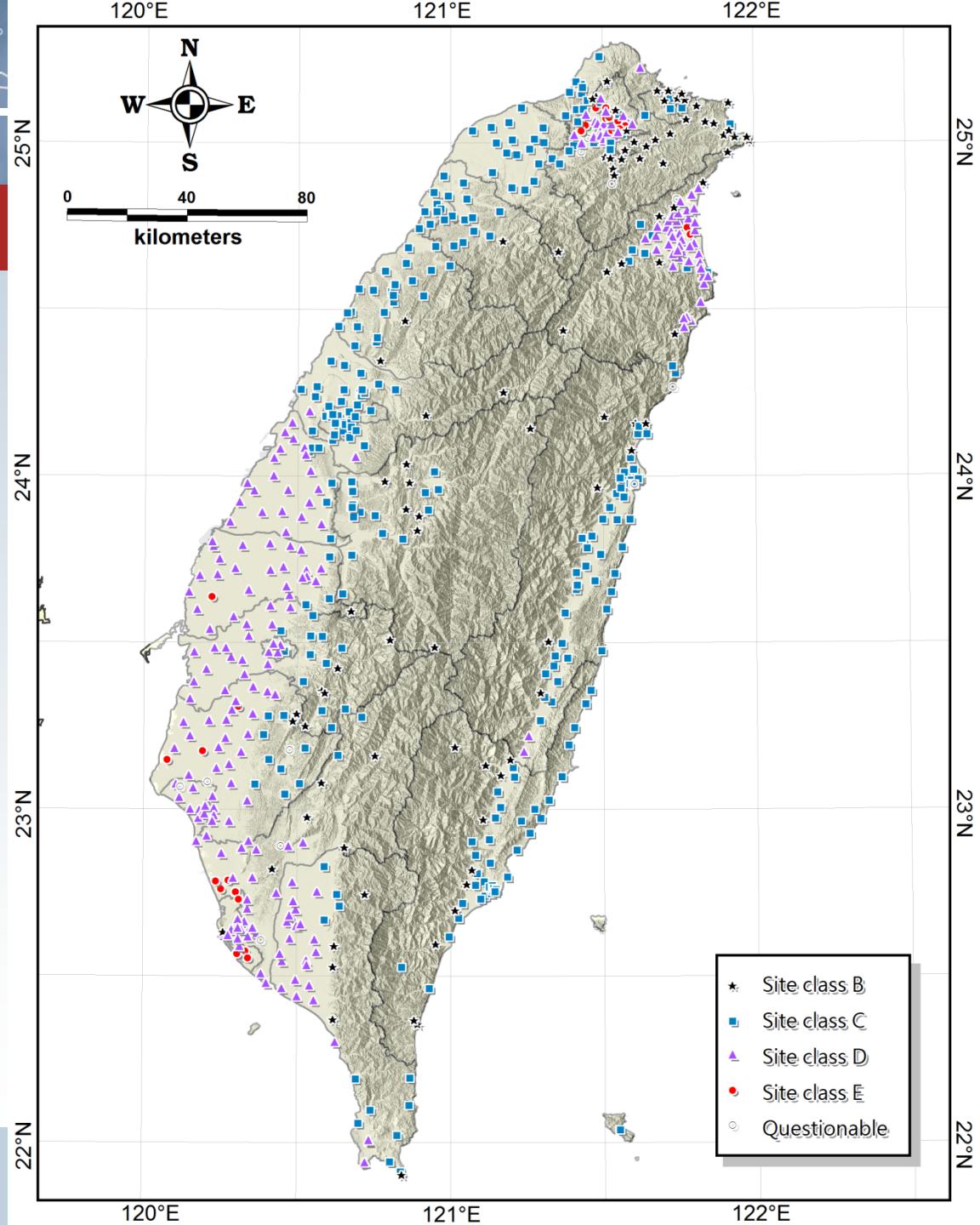
(1900–2007)

- $8 > M_L \geq 7$
- $7 > M_L \geq 6$
- $6 > M_L \geq 5$
- $5 > M_L \geq 4$
- $4 > M_L \geq 3$



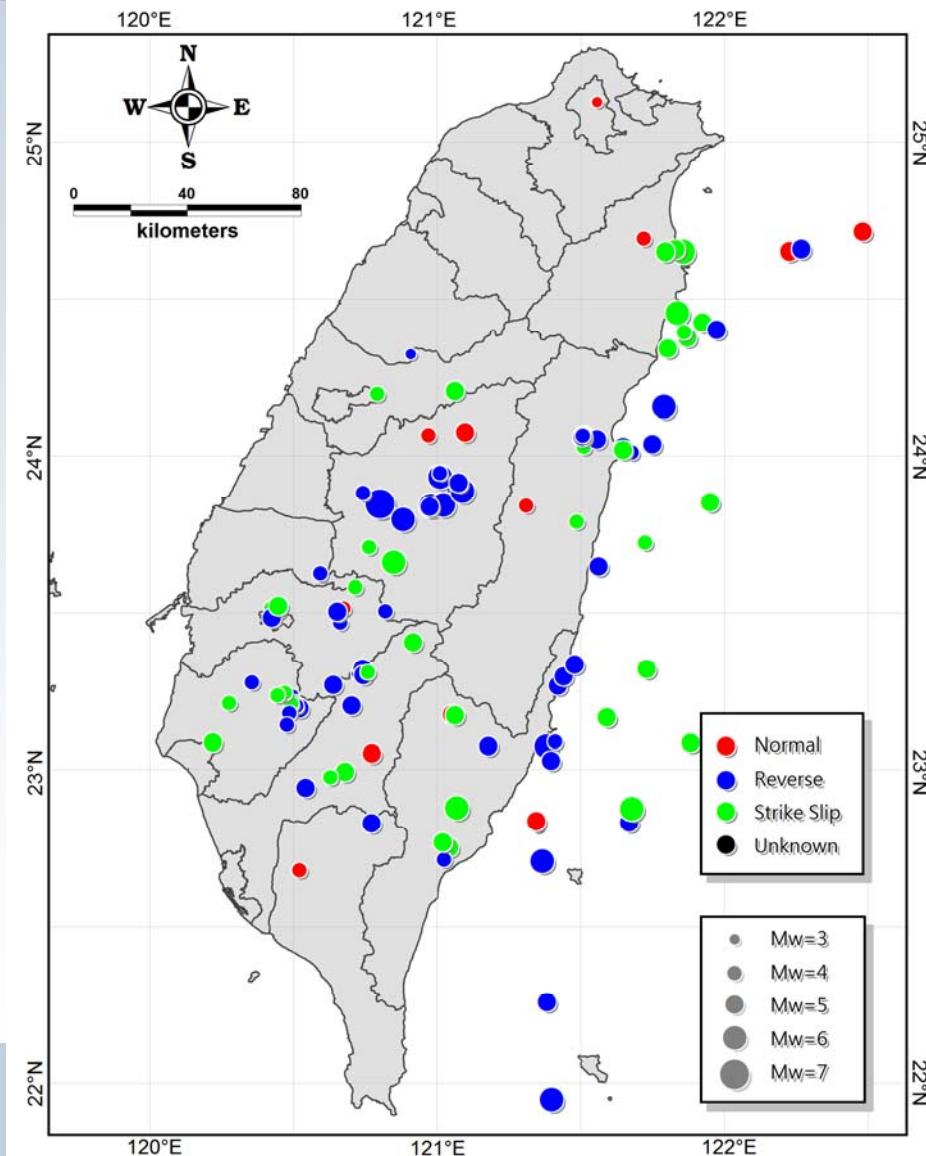
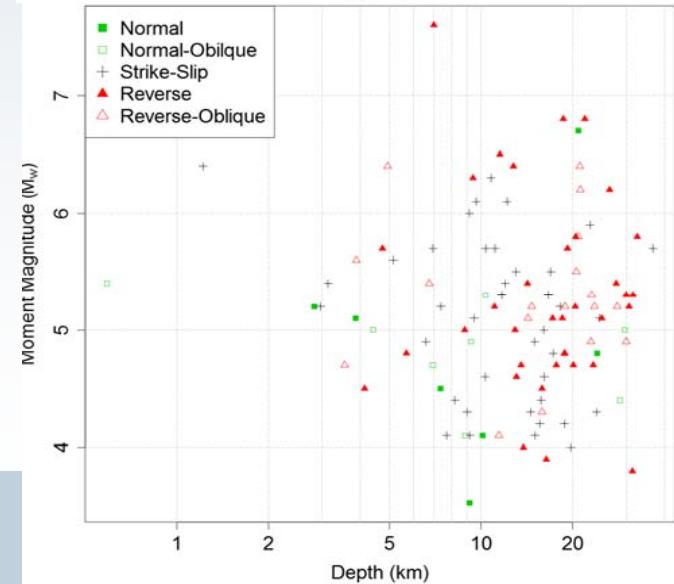
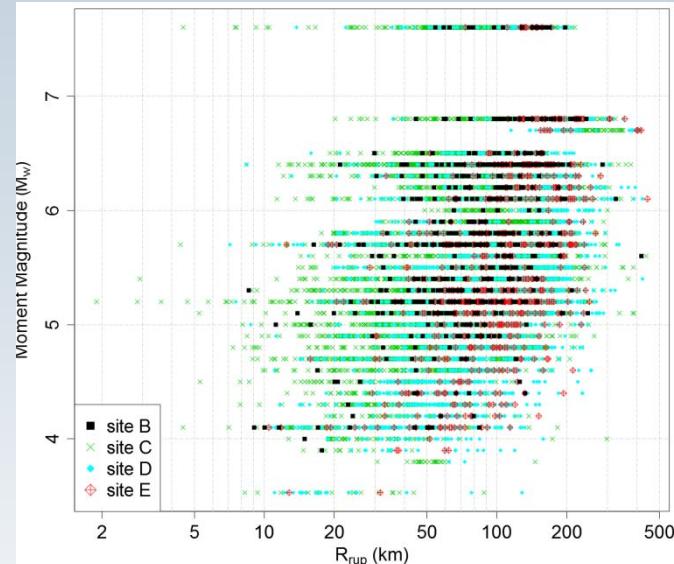
Data source:

- 1.Central Weather Bureau(CWB)
- 2.Institute of Earth Sciences, Academia Sinica



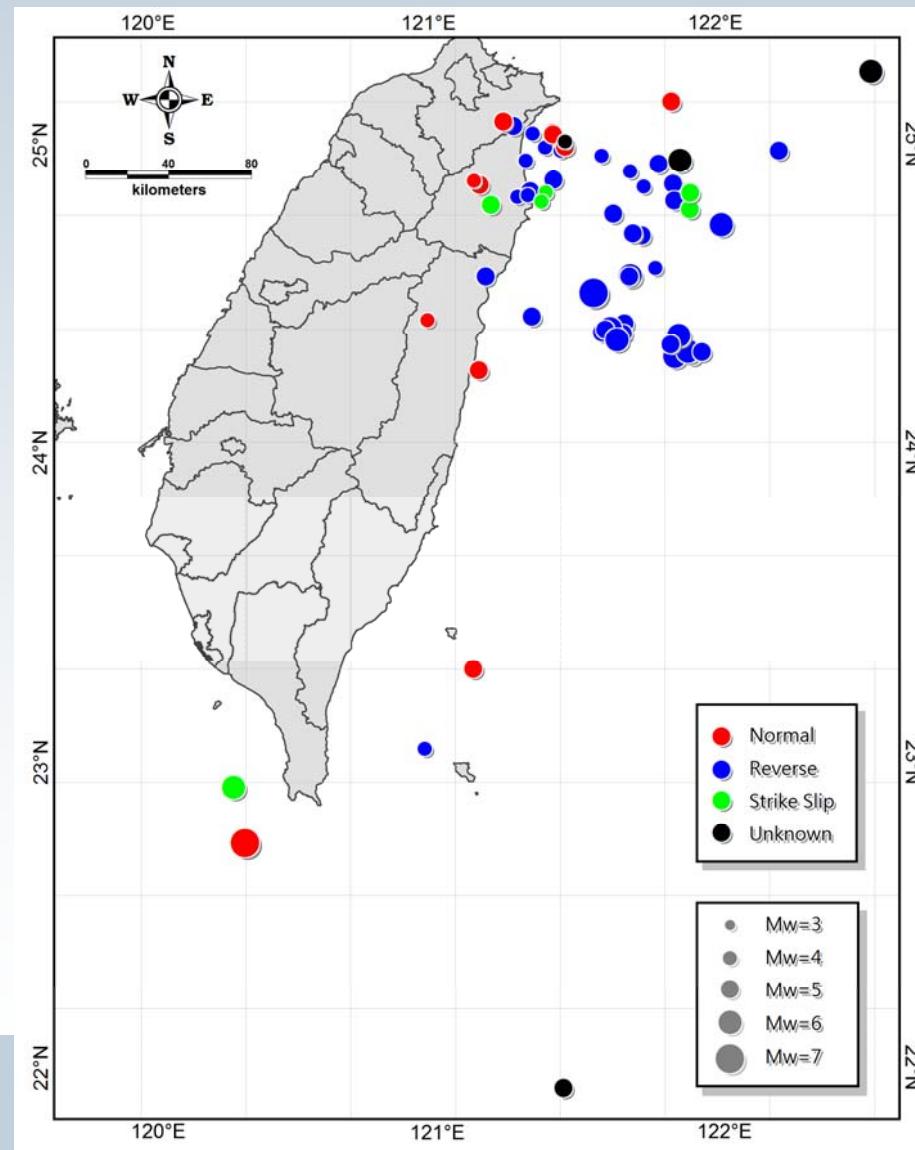
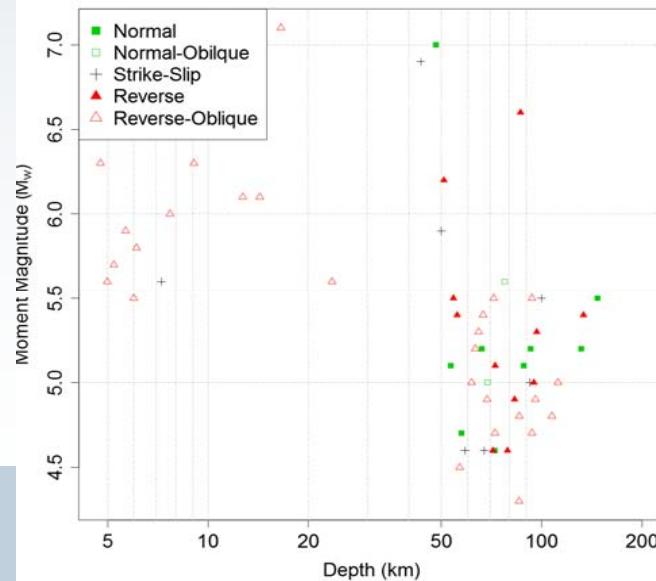
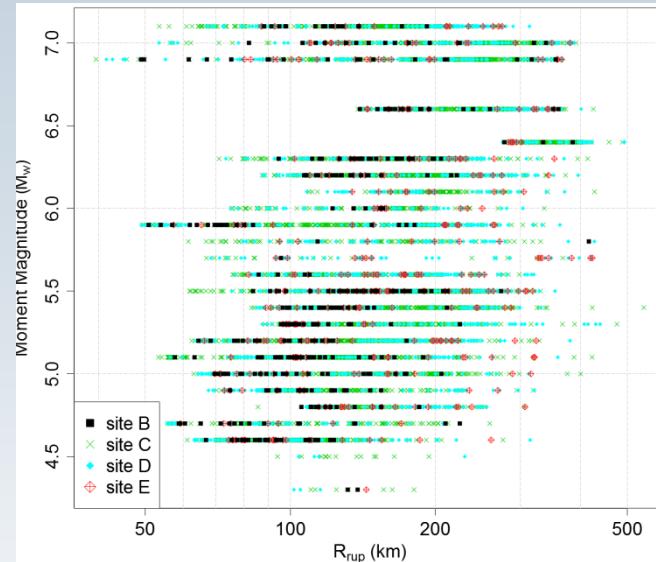


## Strong-motion data for GMPE – Crustal





## Strong-motion data for GMPE – Subduction





# Ground-motion prediction equation

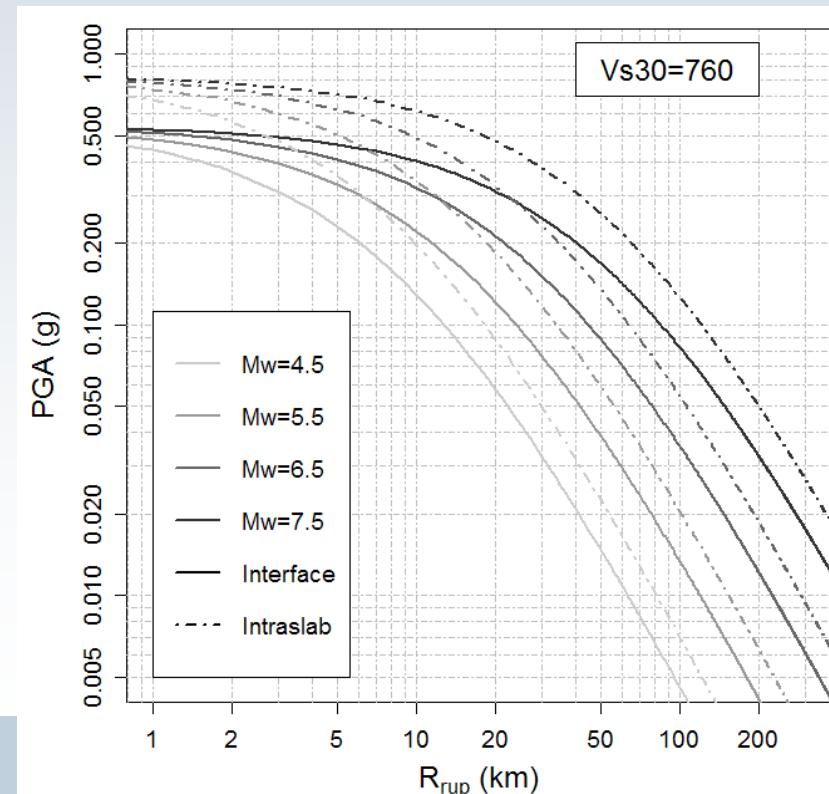
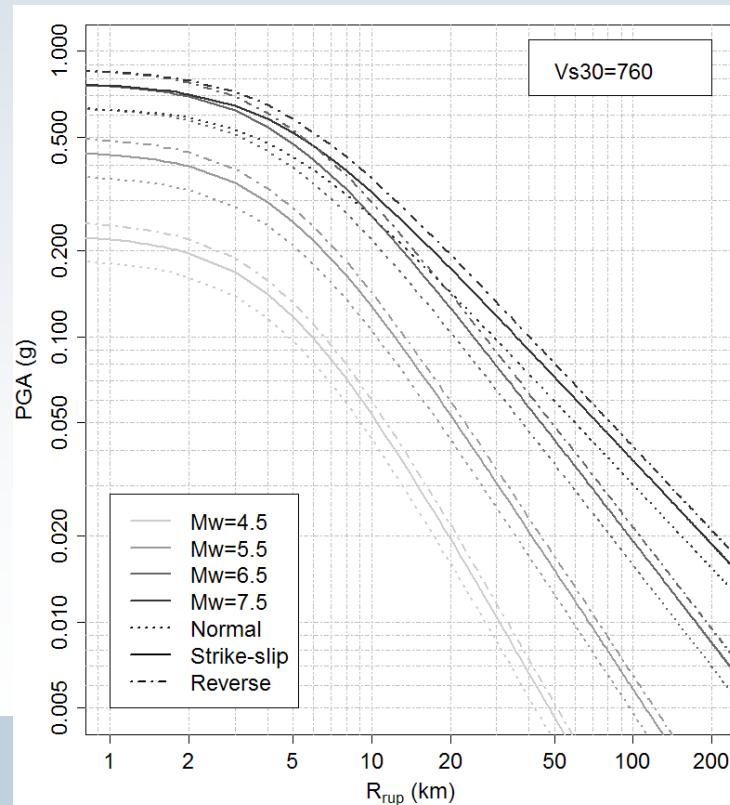
## For Crustal Earthquake

$$\ln y = C_1 + F_1 + C_3(8.5 - M_w)^2 + (C_4 + C_5(M_w - 6.3)) \ln(\sqrt{R^2 + \exp(H)^2}) \\ + C_6 F_{NM} + C_7 F_{RV} + C_8 \ln(VS_{30}/1130)$$

$$\begin{cases} F_1 = C_2(M_w - 6.3) & \text{Where } M_w \leq 6.3 \\ F_1 = (-H \cdot C_5)(M_w - 6.3) & \text{Where } M_w > 6.3 \end{cases}$$

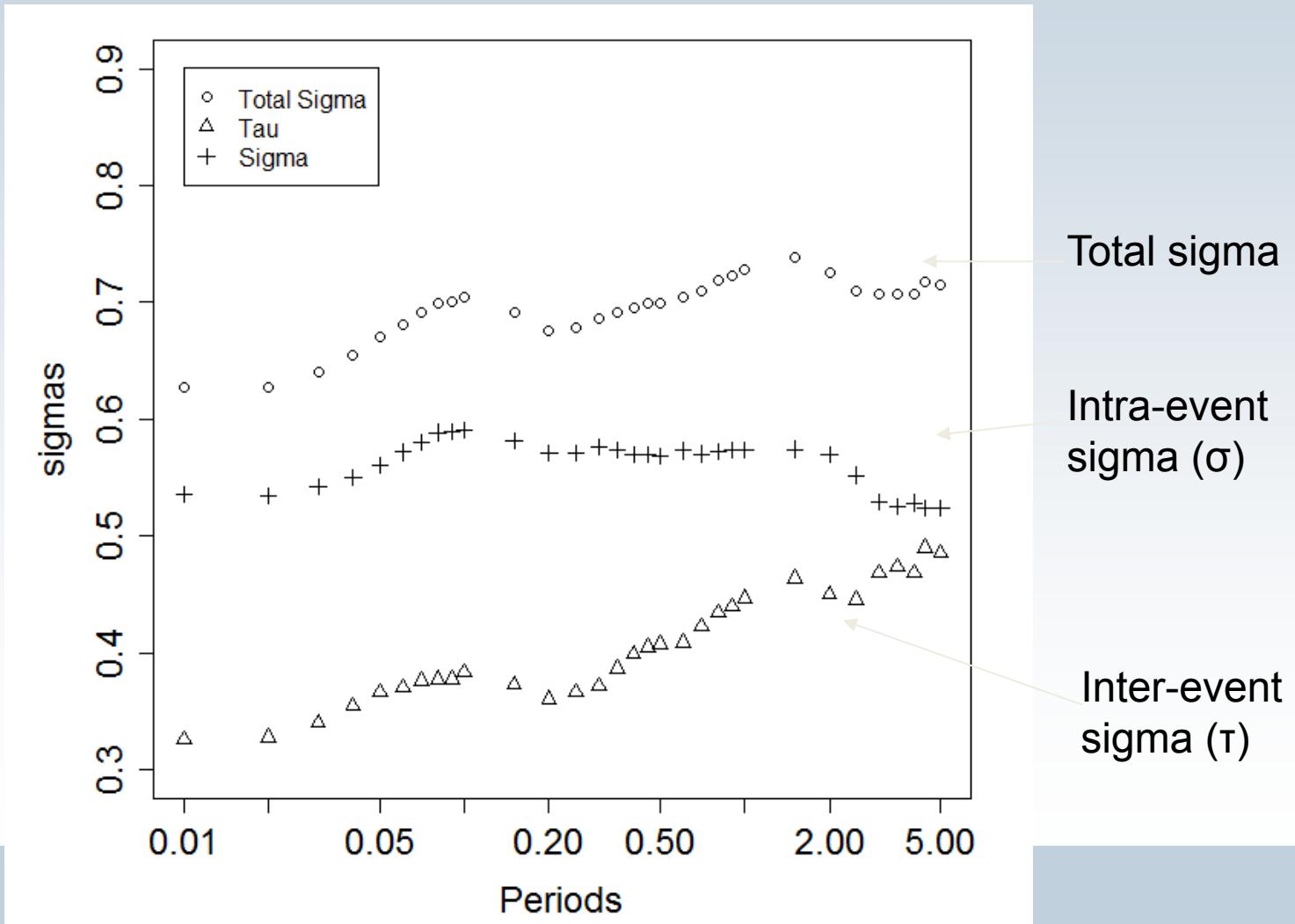
## For Subduction zone earthquake

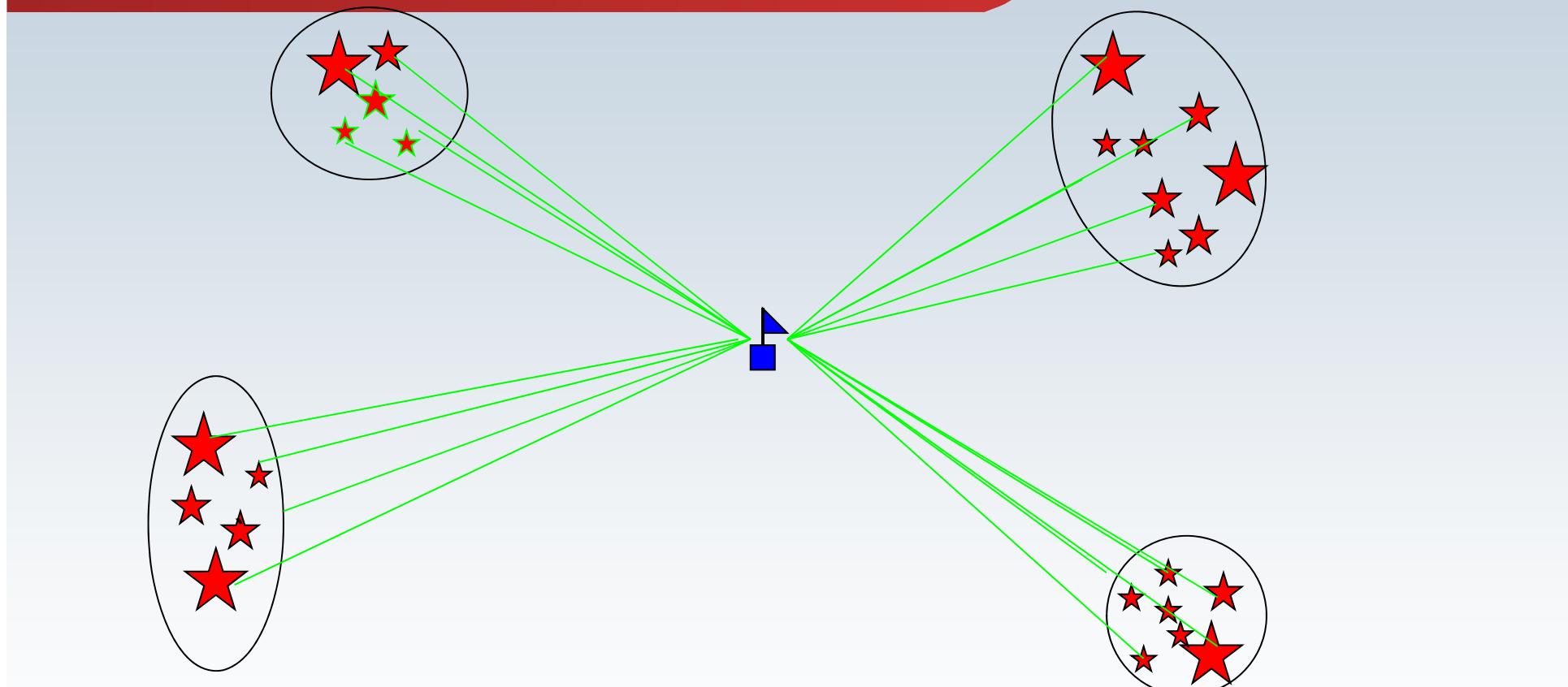
$$\ln y = C_1 + C_2 M + C_3 \ln(R + C_4 e^{C_5 M}) + C_6 H + C_7 Z_t + C_8 \ln(VS_{30}/1130)$$

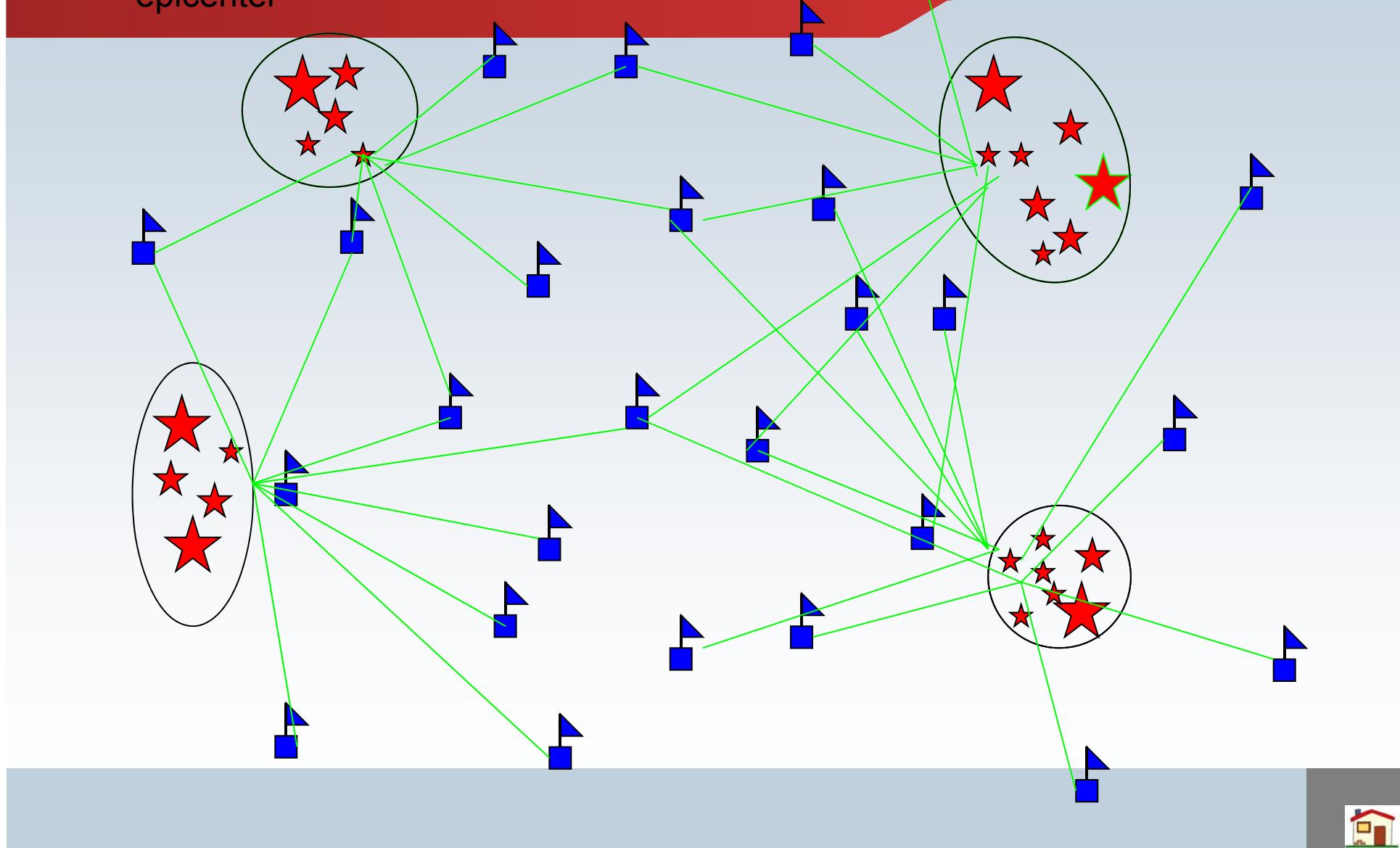




## Variation of the standard deviation in the SA equation with periods from 0.01-5.0 sec







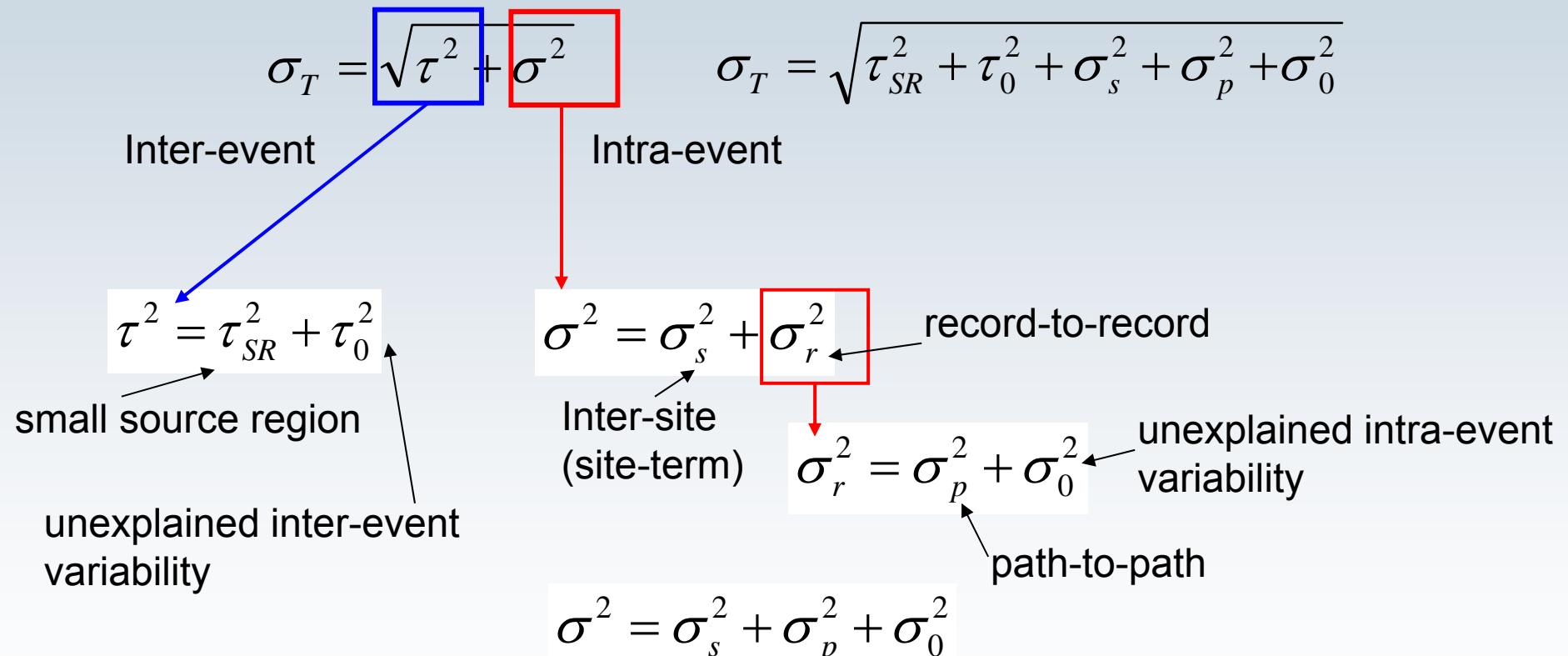


## The notations

- $i \rightarrow$  event
  - Intra-event,  $\xi_{ik}$  inter-event  $\eta_i$
  - Inter-station,  $\xi_{sk}$  intra-station  $\xi_{rik}$
  - Inter-path,  $\xi_{pkl}$  intra-path  $\xi_{0ik}$
  - Inter-region,  $\eta_{SR_l}$  intra-region  $\eta_{0_i}$
- Standard deviations
  - Intra-event,  $\sigma$  inter-event  $\tau$
  - Inter-station,  $\sigma_s$  intra-station  $\sigma_r$
  - Inter-path,  $\sigma_p$  intra-path  $\sigma_0$
  - Inter-region,  $\tau_{SR}$  intra-region  $\tau_0$

Following the notation  
of Walling (2009)

## Decomposition of the variability of empirical ground motion prediction equation - conclusion



Single-path sigma  $\sigma_{sp} = \sqrt{\tau_0^2 + \sigma_0^2}$

Single-site sigma (single-station sigma)  $\sigma_{ss} = \sqrt{\tau^2 + \sigma_r^2}$



# Reduction of the inter-event, intra-event, and total standard deviation from the TSMIP data, if the ergodic assumption is removed

	Single Site		Single Path		
	Total	Intra-event	Total	Inter-Event	Intra-Event
	$\sigma_{SS}$	$\sigma_r$	$\sigma_{SP}$	$\tau_0$	$\sigma_0$
PGA	$0.91\sigma_T$	$0.86\sigma$	$0.54\sigma_T$	$0.69\tau$	$0.43\sigma$
T=0.1	$0.88\sigma_T$	$0.81\sigma$	$0.53\sigma_T$	$0.69\tau$	$0.42\sigma$
T=0.3	$0.90\sigma_T$	$0.86\sigma$	$0.60\sigma_T$	$0.69\tau$	$0.55\sigma$
T=0.5	$0.89\sigma_T$	$0.83\sigma$	$0.61\sigma_T$	$0.69\tau$	$0.57\sigma$
T=1.0	$0.86\sigma_T$	$0.75\sigma$	$0.59\sigma_T$	$0.69\tau$	$0.51\sigma$
T=3.0	$0.86\sigma_T$	$0.69\sigma$	$0.60\sigma_T$	$0.69\tau$	$0.51\sigma$

# Correlation of notation differences for components of variability from previous studies with the current study

Standard Deviation	This Study	Chen & Tsai (2002)	Atkinson (2006)	Morikawa et al (2008)
Total	$\sigma_T$		$\sigma_{\text{reg}}$	$\delta$
Inter-event	$\tau$	$\sigma_E$		$\tau$ (no correction)
Intra-event	$\sigma$			$\sigma$ (no correction)
Inter-site	$\sigma_s$	$\sigma_s$		
Single site, record-to-record	$\sigma_r$	$\sigma_r$		
Inter-path	$\sigma_p$			
Intra-event, Single Path	$\sigma_0$			$\sigma$ (applied correction)
Inter-source region	$\tau_{\text{SR}}$			
Inter-event, Single Region	$\tau_0$			$\tau$ (applied correction)
Single Site (total)	$\sigma_{\text{SS}}$		$\sigma_i$	
Single Path (total)	$\sigma_{\text{SP}}$		$\sigma_{ie}$	(applied correction)

Debris Flow      Climate Change  
Landslide Hazard      Seismic Hazard

中興社防災科技研究中心

# Comparison of single-path standard deviations ( $\sigma_{SP}$ ) as a fraction of $\sigma_T$

	This Study	Atkinson (2006)	Morikawa et al. (2008)
PGA	$0.54\sigma_T$	$0.67\sigma_T$	$0.46\sigma_T$
T=0.1	$0.53\sigma_T$		$0.38\sigma_T$
T=0.3	$0.60\sigma_T$	$0.68\sigma_T$	$0.44\sigma_T$
T=0.5	$0.61\sigma_T$		$0.45\sigma_T$
T=1.0	$0.59\sigma_T$	$0.67\sigma_T$	$0.47\sigma_T$
T=3.0	$0.60\sigma_T$		$0.47\sigma_T$



## Conclusion

- More and more strong-motion data
- Strong ground motion parameters
  - PGA, PGV, PGD, Ia and Sa
- Including crustal earthquake and subduction zone earthquake
- Source-to-site distance
  - From epicenter distance to the closest distance to fault rupture plane
- Site parameters
  - From rock/soil to Vs30
- Hanging-wall and footwall



## Future work

- For GMPE
  - More predictor variables for source, path and site
  - Directivity
  - Nonlinear site effect
  - Strong-motion difference due to various stress drop
- For Ground-motion prediction
  - Empirical approach
  - Seismological approach
    - Empirical Green's functions (EGFs)
    - Stochastic Green's functions (SGFs)
    - Hybrid Green's functions (HGFs)



Thank you for your attention.