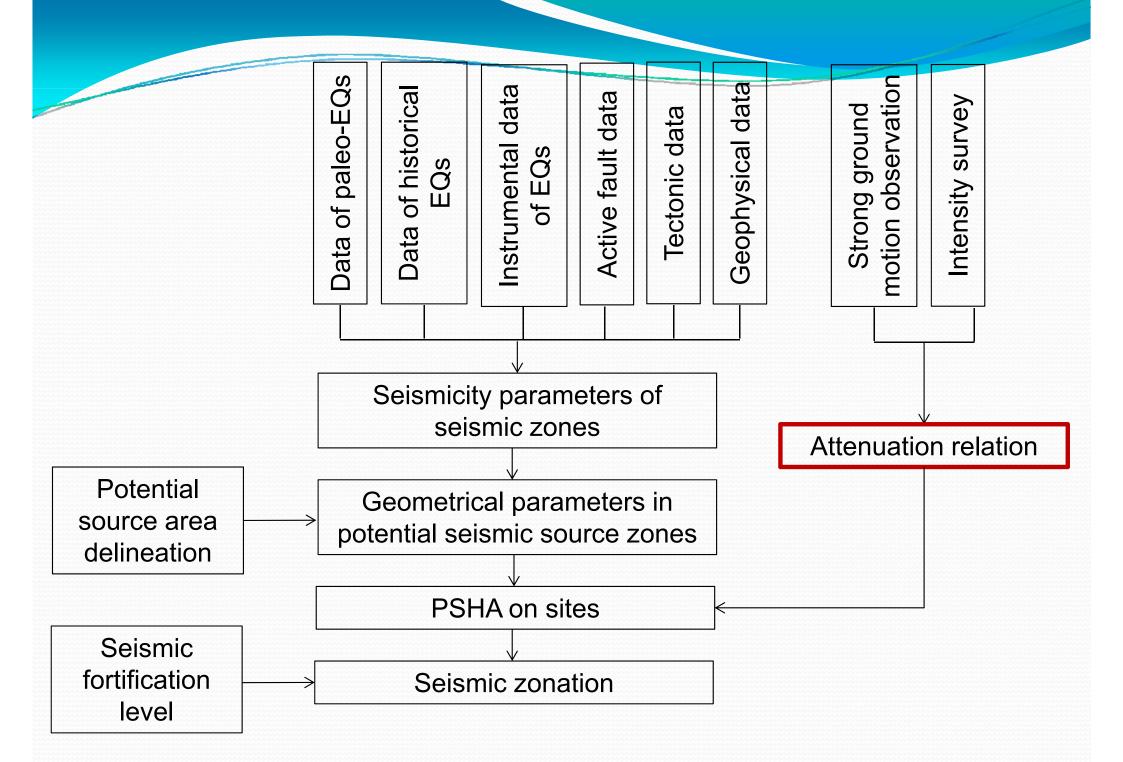
Seismology based ground motion attenuation relationship for the demonstration region of China in the joint project

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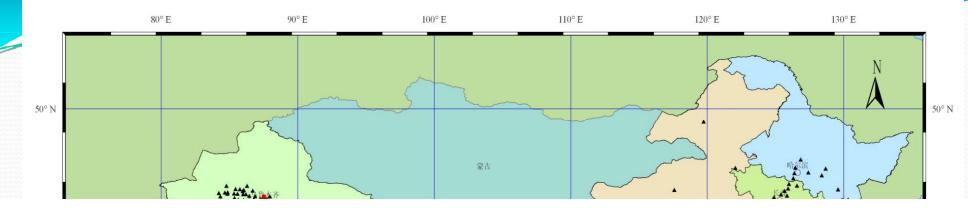
2013.6.17, Sendai, Japan

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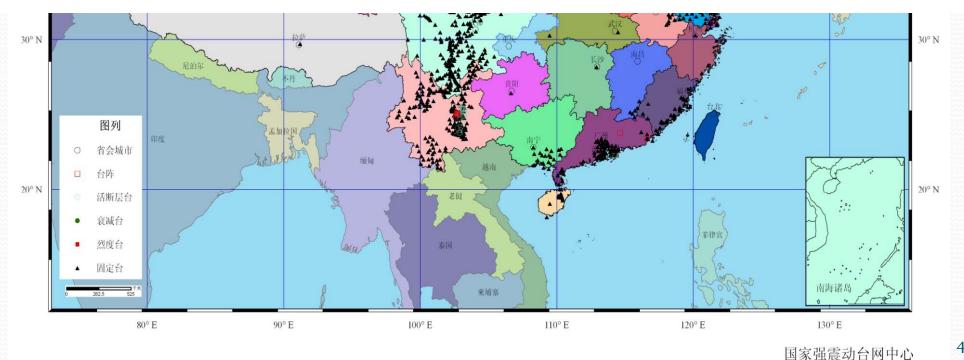
- Background
- Methodology
- Demonstration region
- Sichuan-Yunnan region



#### 中国数字强震动台网分布图



#### How to establish relationships for regions with few or without strong ground motion records?



- In China, strong ground motion records are not enough now.
- Mapping method (Hu, 1980s):

This method assumes that, for region A of enough acceleration observation data and a region B of few such data, earthquake pairs ( $M_A$ ,  $R_A$ ;  $M_B$ ,  $R_B$ ) exist in the intensity attenuation curves  $I_A(M_A, R_A)$  of region A and  $I_B(M_B, R_B)$  of region B, so that they give the same intensity I and ground motion Y.

$$P(M_A, R_A) \rightarrow [I_A(M_A, R_A) = I_B(M_B, R_B)] \rightarrow Q(M_B, R_B)$$

 $\ln Y_{a}(M_{A}, R_{A}) \rightarrow \square \rightarrow \ln Y_{B}(M_{B}, R_{B})$ 

Western US:

 $I = 0.514 + 1.500 M - 0.00659 R - 2.014 \log(R + 10), \sigma = 0.274, R < 300 km$  $\log(PGA) = 1.297 + 0.566 M - 1.723 \log[R + 1.046e^{0.451M}], \sigma = 0.240$ 

Eastern China: 
$$\begin{cases} I_a = 5.019 + 1.446M - 4.1361g(R + 24) \\ I_b = 2.240 + 1.446M - 3.0701g(R + 9) \end{cases}, \quad \sigma = 0.517 \end{cases}$$

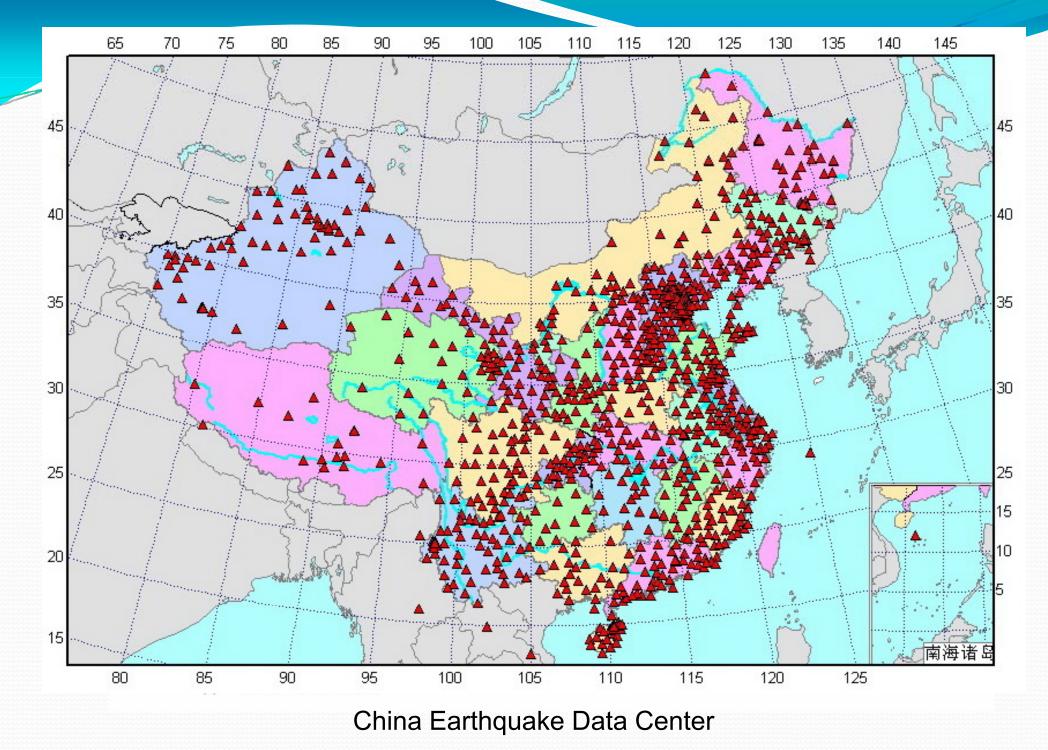
Western China: 
$$\begin{cases} I_a = 5.253 + 1.398M - 4.164 \lg(R+26) \\ I_b = 2.019 + 1.398M - 2.943 \lg(R+8) \end{cases}, \quad \sigma = 0.632 \end{cases}$$

#### Eastern China:

$$\begin{cases} \log(PGA)_a = 2.027 + 0.548M - 1.902 \log[R + 1.700e^{0.425M}] \\ \log(PGA)_b = 1.035 + 0.519M - 1.465 \log[R + 0.381e^{0.525M}] \end{cases}, \quad \sigma = 0.240$$

Western China:

$$\begin{cases} \log(PGA)_a = 2.206 + 0.532M - 1.954 \log[R + 2.018e^{0.406M}] \\ \log(PGA)_b = 1.010 + 0.501M - 1.441 \log[R + 0.340e^{0.521M}] \end{cases}, \quad \sigma = 0.240$$



## Methodology

Source Fourier spectrum of ground motion  $FA(M_0, f, R)$  from a point source can be described as

 $FA(M_0, f, R) = C \cdot S(M_0, f) \cdot G(R) \cdot D(R, f) \cdot A(f) \cdot P(f) \cdot I(f)$ 

where, *C* is proportion factor;  $S(M_0, f)$  is source spectrum for a specified seismic moment; G(R) is geometric spreading function; D(R, f) is anelastic attenuation function; A(f) is the amplification factor of near surface amplitude; P(f) is a high-cut filter that rapidly reduces amplitudes at high frequencies; I(f) is spectrum shape parameter, used to shape the spectrum to correspond to the particular ground-motion measure of interest.

 $C = \frac{R_{\theta\phi}FV}{4\pi R_0 \rho_s \beta_s^3}$ 

$$S(M_0, f) = \frac{M_0}{1 + \left(\frac{f}{f_0}\right)^2}$$

$$f_0 = 4.9 \times 10^6 \,\beta (\Delta \sigma \,/\,M_0)^{1/3}$$

$$G(R) = \begin{cases} \frac{1}{R} & R \leq R_1 \\ \frac{1}{R_1} & R_1 < R < R_2 \\ \frac{1}{R_1} \sqrt{\frac{R_2}{R}} & R \geq R_2 \end{cases}$$

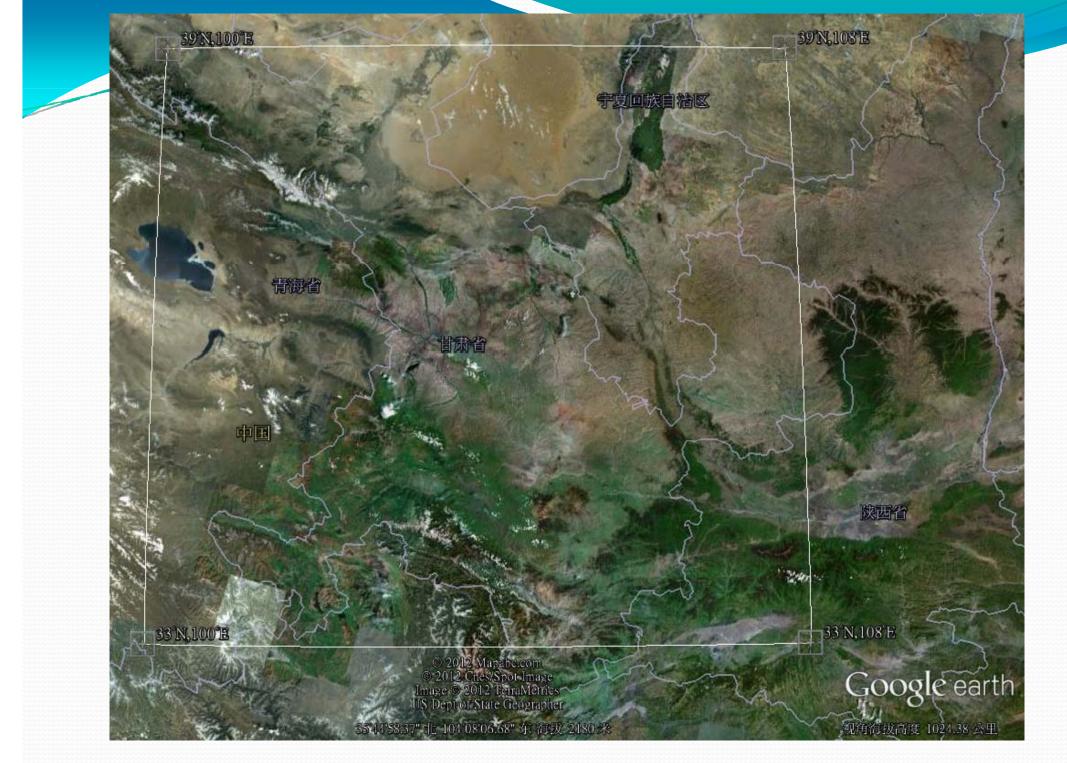
$$D(R,f) = \exp\left(-\frac{\pi fR}{Q\beta}\right)$$

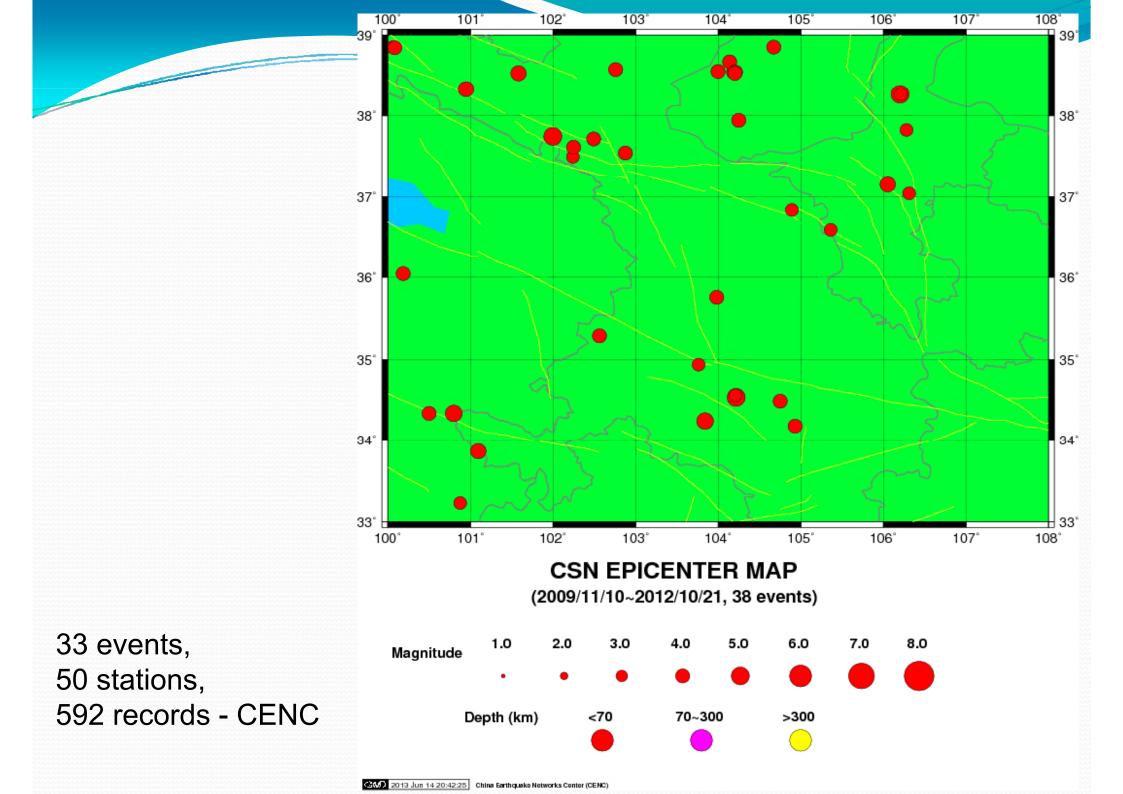
$$Q = Q_0 f^{\eta}$$

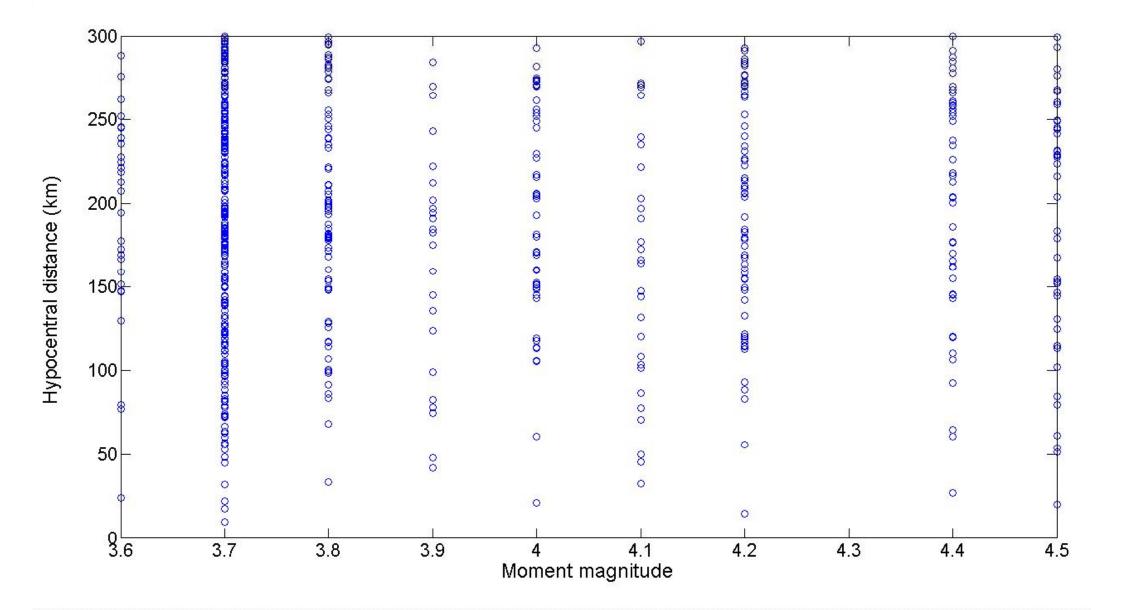
$$P(f) = \left[1 + \left(\frac{f}{f_{\text{max}}}\right)^8\right]^{-1/2}$$

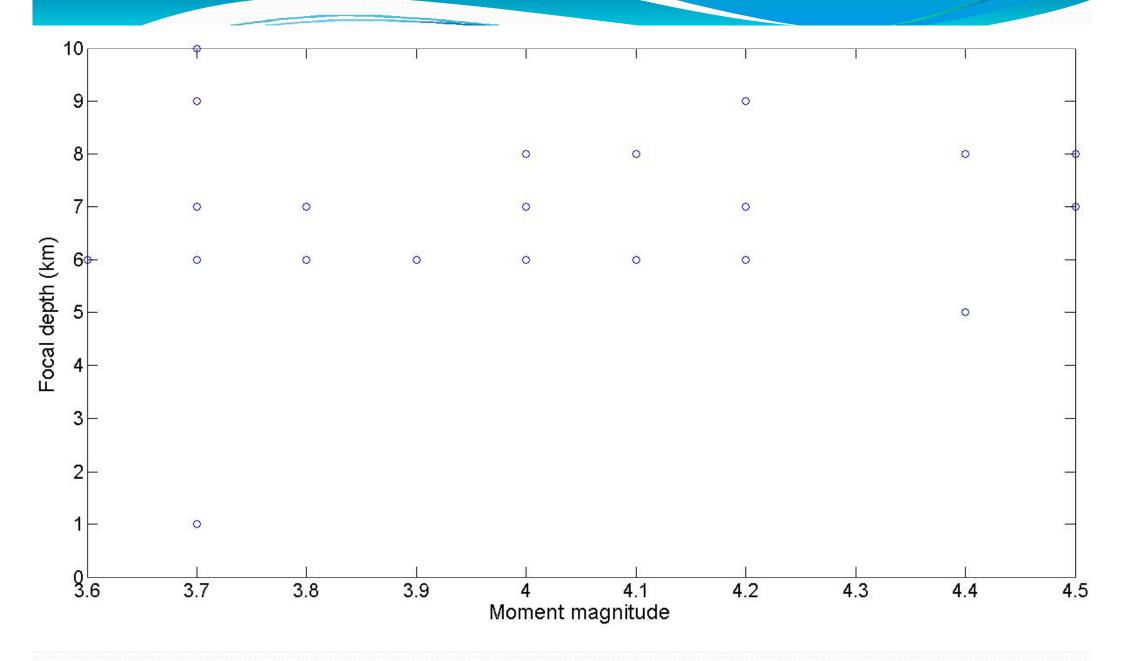
$$I(f) = (2\pi f)^z$$

#### No available values









## Inversion strategy

- µGA is developed from GA, the main procedure is similar.
- In µGA, there still retain selection and crossover, but no variation, at the same time, it retains the optimal individual.

Objective function

$$\phi_{j} = \sum_{m} \sum_{n} [FA_{0}(m,n) - FA_{j}(m,n)]^{2}$$

Fitness

$$F_j = e^{-\beta\phi}$$

#### Inversion ranges

-

Δσ (bars)	Q <sub>o</sub>	η	R <sub>1</sub> (km)	R <sub>2</sub> (km)	
1~200	200~700	0.2~0.6	50~100	100~150	

#### Inversion results

Δσ (bars)	Q <sub>o</sub>	η	R <sub>1</sub> (km)	R <sub>2</sub> (km)	
90.99	324.98	0.39	65.30	119.28	

### **Time-domain simulation**

- Combined with random phase spectra, the Fourier spectra can be transformed into the time domain.
- In time-domain simulations, these time series are windowed by

J

$$f(t) = \begin{cases} (t/t_1)^2, & 0 \le t \le t_1 \\ 1.0, & t_1 < t \le t_2 \\ e^{-c(t-t_2)}, & t_2 < t \end{cases}$$

where,  $t_1$  and  $t_2$  are the starting point and finishing point of the stable section, c is the attenuation rate.

An envelop curve on bedrock (Huo and Hu, 1991) is adopted.

$$\begin{cases} \lg t_1 = -1.074 + 1.005 \lg (R + 10) \\ \lg t_2 = -2.268 + 0.3262 M_w + 0.5815 \lg (R + 10) \\ \lg c = 1.941 - 0.2817 M_w - 0.5870 \lg (R + 10) \end{cases}$$

- The windowed time histories are transformed into the frequency domain.
- Complex spectra are transformed back to the time domain.
- Attenuation curves, which show mean levels, can be constructed after mean PGAs from 50 different random phase spectra are calculated.

## Strong motion data from CSMNC

- 1987.9.12-2011.11.2
- 18 events
- 21 stations
- Mw=5.0, 33 records
- Mw=6.0, 5 records
- Mw=7.0, 1 record

#### **Empirical relations**

1989-2012 Seven sets of relations-empirical Western China, Loess region of China

- Zoning map, Western China (2001)  $\ln(PGA) = 5.912 + 1.836M - 2.846 \ln(R + 3.400 \cdot e^{0.451M})$  $\ln(PGA) = 2.509 + 1.36M - 1.791 \ln(R + 1.046 \cdot e^{0.451M})$
- Model 1 (2012)

 $\log_{10}(PGA) = 0.617 + 1.163M - 0.046M^2 - 2.207 \log_{10}(R + 1.694 \cdot e^{0.446M})$ 

 $\log_{10}(PGA) = -0.644 + 1.08M - 0.043M^2 - 1.626\log_{10}(R + 0.255 \cdot e^{0.57M})$ 

#### XIAO Liang (2010)

 $\log_{10}(PGA) = 1.204 + 0.632M - 1.928 \log_{10}(R + 1.046 \cdot e^{0.451M})$ 

YU Yanxiang and WANG Suyun (Western China, 2006)

 $\log_{10}(PGA) = 2.206 + 0.532M - 1.954 \log_{10}(R + 2.018 \cdot e^{0.406M})$ 

 $\log_{10}(PGA) = 1.01 + 0.501M - 1.441\log_{10}(R + 0.340 \cdot e^{0.521M})$ 

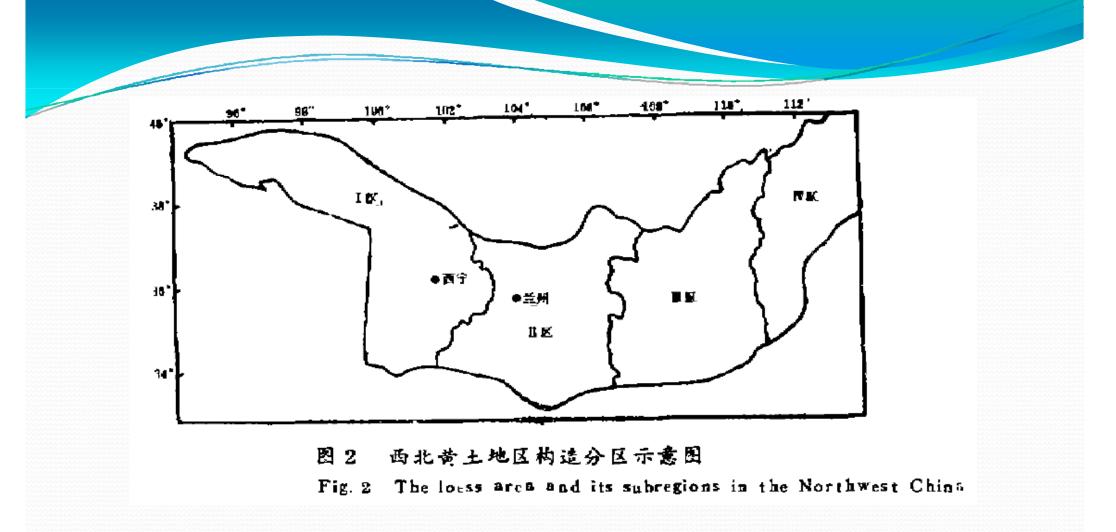
 DING Boyang, TIAN Shaobo and LEI Zhongsheng (Loess region of China, 1991). Circle model, IR and IM

 $\ln(PGA) = 1.066 + 0.559M - 2.012\ln(R + 17) - 0.00387(R + 17)$ 

 $\ln(PGA) = 0.071 + 0.500M - 1.502\ln(R + 17)$ 

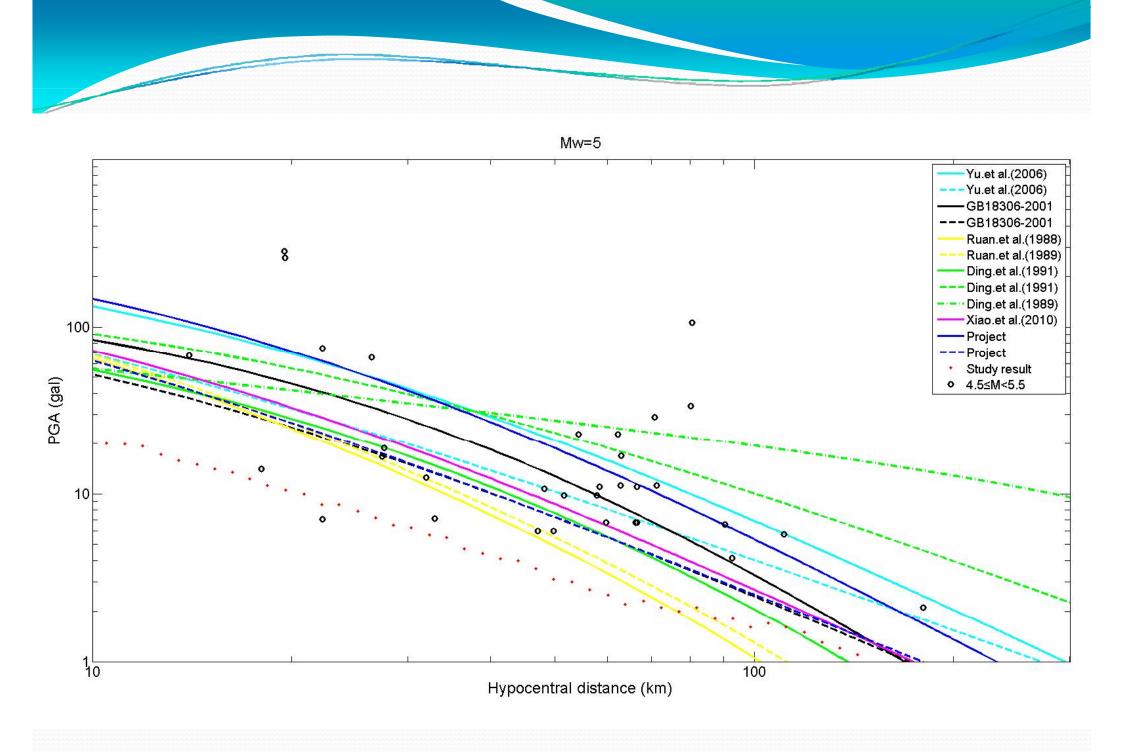
 DING Boyang, LEI Zhongsheng and FANG Shulan (Loess region of China, 1989). Circle model

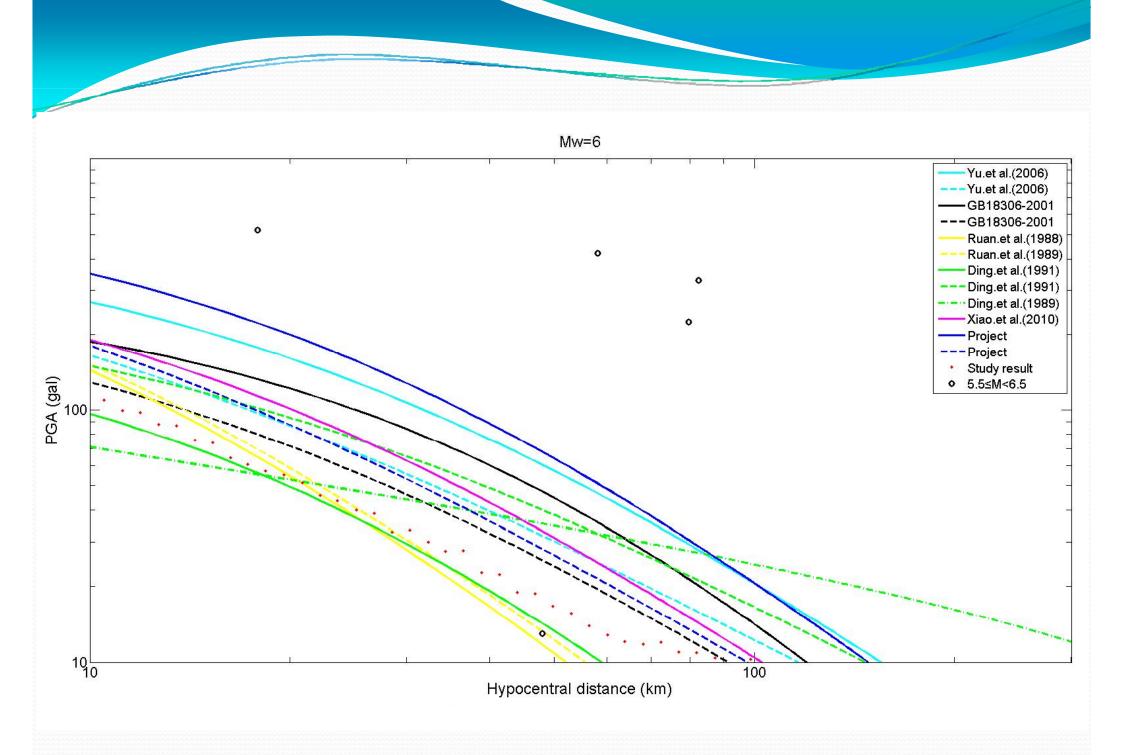
 $\ln(PGA) = -3.066 + 0.2347M - 0.4137 \ln R - 0.00127R$ 

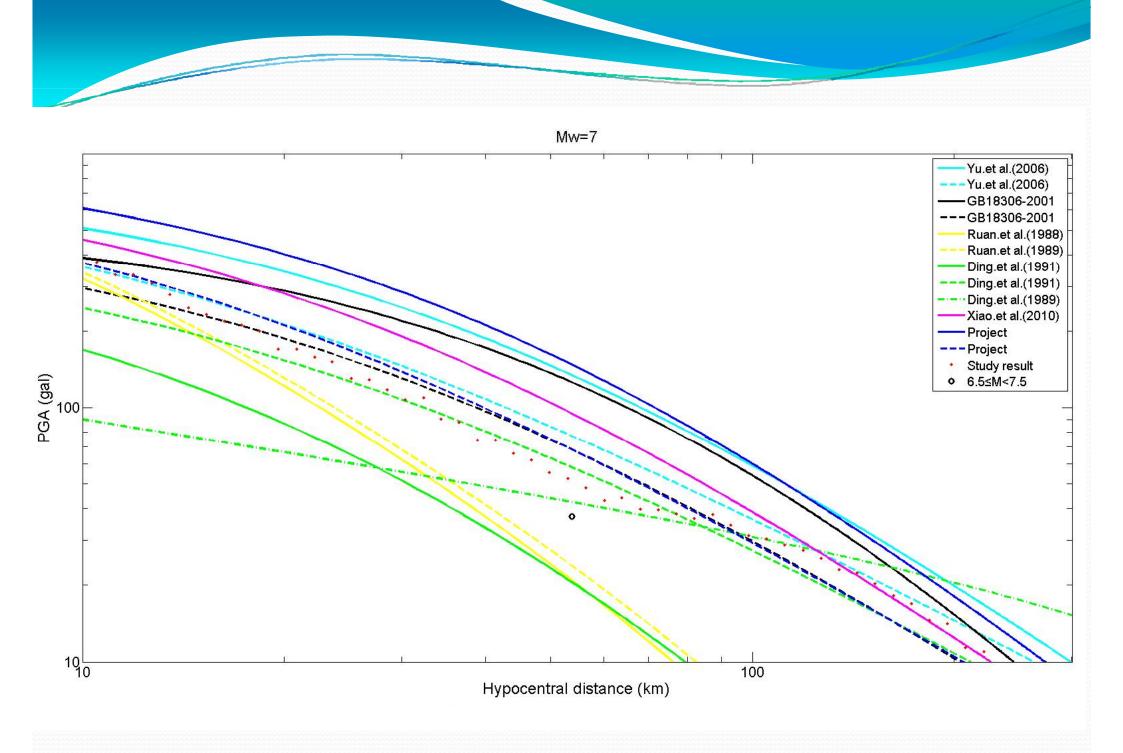


RUAN Aiguo and SUN Congshao, A study on attenuation law of seismic ground motion in the loess region of the Northwest China, 1989

 $\ln(PGA) = -2.9169 + 0.8028 M - 1.508 \ln \sqrt{R^2 + 36} - 0.00985 \sqrt{R^2 + 36}$  $\ln(PGA) = -2.888 + 0.8038 M - 1.507 \ln \sqrt{R^2 + 36} - 0.00815 \sqrt{R^2 + 36}$ 

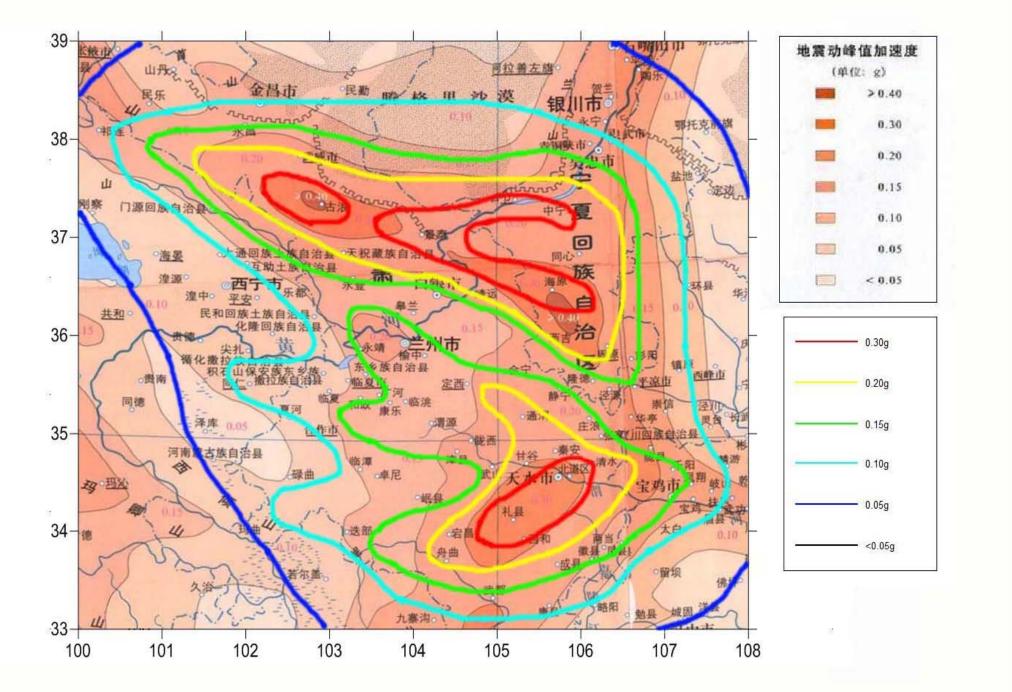






## **Application in PSHA**

R	20.9	22.2	23.6	25.2	26.8	28.5	30.3	32.3	34.3	36.5
5.5	24.8	21.9	20.2	19.2	17.3	16.6	15.2	14.4	12.9	12.1
5.6	29	24.8	24.3	22.4	21.1	18.4	18.6	16.6	15.3	14.9
5.7	32.7	29.5	29.2	25.3	25.2	22.6	20.5	20.6	17.7	15.8
5.8	38.4	35.6	32.9	32.3	27.5	26.5	23.3	22.2	20.3	19.6
5.9	43.5	40.2	36.7	35.3	31.9	31.7	27	26.3	23.6	21.8
6.0	51.4	45.8	43.7	40	38.8	32.9	33.5	30	27.6	27.8



## Conclusions

- In the demonstration region of this project, five regional parameters are inversed from small earthquakes by µGA.
- Strong ground motion attenuation relationship is constructed from small earthquake records by China Earthquake Network.
- The relation is expressed as a 2D table and is adopted in PSHA.
- The seismic zoning map of the demonstration is compared with the national map to prove the validation.

# The end

# Thanks